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Hydroelectric Dams Planned on Brazil's Xingu River as sources of Greenhouse Gases: Belo Monte (Kararaô) and Altamira (Babaquara)

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Abstract Estimating the greenhouse-gas emissions from hydroelectric dams is important as an input to the decision-making process on public investments in the various options for electricity generation and conservation. Brazil's proposed Belo Monte Dam (formerly Kararaô) and its upstream counterpart, the Altamira Dam (better known by its former name: Babaquara) are at the center of controversies regarding how greenhouse-gas emissions from dams should be counted. The Belo Monte Dam by itself would have a small reservoir area (440 km²) and large installed capacity (11,181.3 MW), but the Babaquara Dam that would regulate the flow of the Xingu River (thereby increasing power generation at Belo Monte) would flood a vast area (6140 km²). The water level in Babaquara would rise and fall by 23 m each year, annually exposing a drawdown area of 3580 km² on which soft easily decomposed vegetation would quickly grow. This vegetation would decompose each year at the bottom of the reservoir when the water level rises, producing methane. The methane from drawdown-zone vegetation represents a permanent source of this greenhouse gas, unlike the large peak of emission from decomposition of initial stocks of carbon in the soil and in the leaves and litter of the original forest. The turbines and spillways draw water from below the reservoir's thermocline, releasing a large part of the dissolved methane to the atmosphere. Carbon dioxide from decay of the above-water portions of trees in the forest that is flooded represents another significant greenhouse gas emission source in the early years after reservoir formation. Belo Monte and Babaquara represent a challenge to Brazil's

fledgling environmental impact assessment and licensing system because the current procedure considers only one infrastructure project at a time, rather than the full range of impacts that the overall development entails. In this case, the unusually favorable characteristics of the first dam (Belo Monte) are highly misleading as indications of the environmental consequences of a decision to build that dam: the principal impacts are in the impetus it gives to creating much larger upstream reservoirs, beginning with Babaquara and possibly including an additional four dams planned for the Xingu Basin, all of which flood tropical rainforest and indigenous land in addition to emitting greenhouse gases. The present analysis indicates that the Belo Monte/Babaquara complex would not break even in terms of greenhouse gas emissions until 41 years after the first dam is filled in a calculation with no discounting, and that any annual discount rate above 1.5% results in the complex failing to perform as well as natural gas by the end of the 50-year time horizon used in Brazil's assessments of proposed energy projects. The global-warming impact of dams is one indication of the need for Brazil to reassess its current policies that allocate large amounts of energy in the country's national grid for aluminum smelting for a subsidized export industry.

Amazonia – Brazil – Global warming - Greenhouse-gas emissions - Hydroelectric dams
- Methane - Reservoirs

The proposed Belo Monte Dam, on Brazil's Xingu River (a north-flowing tributary to the Amazon in the State of Pará) is the focus of intense controversy due to the magnitude and nature of its impacts. Belo Monte's impact on global warming stems from the upstream dams that would add substantially to Belo Monte's electrical output by regulating the flow of the highly seasonal Xingu River. The Belo Monte reservoir itself is small relative to the capacity of its two powerhouses, but the five upstream reservoirs are vast, even by Amazonian standards. The largest of these is the Babaquara Dam--recently re-named the "Altamira Dam" in an apparent effort to escape the onus of the criticism that the plans for Babaquara have attracted over the past two decades (the initial survey, or *inventário*, began in October 1975).

ELETRONORTE (the government power authority in northern Brazil) first proposed the Kararaô Dam (now called "Belo Monte") with power generation calculations that assumed upstream regulation of streamflow by at least one dam (Babaquara)(CNEC 1980). The series of dams on the Xingu River would have serious consequences for indigenous peoples and for the large areas of tropical rainforest that the reservoirs would flood (*e.g.*, Santos and de Andrade 1990, Sevá and Switkes forthcoming). Difficulties in obtaining environmental approval led to formulation of a second plan for Belo Monte, with calculations that assumed no upstream regulation of streamflow (Brazil, ELETRONORTE 2002). The viability study for the second plan makes clear that the need for an analysis under the assumption of unregulated streamflow stems from "a new economic and socio-environmental perspective" (*i.e.*, from political considerations), and that Belo Monte's power output would be greater if upstream dams were included in the analysis (Brazil, ELETRONORTE 2002, p. 6-82). Further difficulties in obtaining environmental approval have led ELETRONORTE to initiate a third analysis (still in progress) with several possible smaller installed capacities: 5500, 5900 and 7500 MW (Pinto 2003). These developments in no way imply that a decision has been made not to build the Babaquara (Altamira) Dam upstream of Belo Monte. On the contrary, preparations for construction of Babaquara

(Altamira) are included under the current Decennial Plan for the electrical sector (Brazil, MME-CCPESE 2002) and plans for the dam are presented by ELETRONORTE as progressing normally towards construction (Santos 2004). In other words, the one-dam scenario portrayed in the Belo Monte viability study (Brazil, ELETRONORTE 2002) and environmental impact study (Brazil, ELETRONORTE nd [C. 2002]a) appears to represent a bureaucratic fiction that was drafted for the purpose of gaining environmental approval for Belo Monte. The scenario used in the present paper therefore appears most likely as a representation of the overall impact of the development, with the Belo Monte Dam being built according to the 2002 viability study (Brazil, ELETRONORTE 2002), followed by Babaquara (Altamira) in accord with earlier plans (Brazil, ELETRONORTE nd [C. 1988]). Belo Monte cannot be considered alone without taking into account the impacts of upstream dams, particularly Babaquara (Altamira). Among the many impacts of the upstream dams that must be assessed, one is the role they would play in creating greenhouse-gas emissions. In the present analysis, preliminary estimates will be presented for emissions from Belo Monte and Babaquara; if the remaining four planned dams were built, they would have additional impacts.

Hydroelectric Dams and Greenhouse-Gas Emissions

Belo Monte is at the heart of ongoing controversies regarding the magnitude of the global-warming impact of hydroelectric dams and the proper way that this impact should be quantified and considered in the decision-making process. When the first calculations of greenhouse-gas emissions from the existing dams in Brazilian Amazonia indicated substantial impact (Fearnside 1995a), these conclusions were attacked by presenting a hypothetical case that corresponded to Belo Monte with a power density of over 10 Watts of installed capacity per m^2 of reservoir surface area (Rosa and others 1996). In addition to methodological issues that make these hypothetical calculations underestimates of greenhouse-gas impact, the main problem is the omission of the emissions from the 6140- km^2 Babaquara Dam upstream of Belo Monte (Fearnside 1996a). This same basic problem remains today, even after many advances in greenhouse-gas emissions estimates.

The relatively small area of the Belo Monte Dam by itself means that greenhouse-gas emissions from the reservoir surface will be modest, and when these are divided by the Dam's 11,181 MW installed capacity the emission appears to be low relative to the benefits. This is the rationale of using the "power density" (Watts of installed capacity per square meter of water area) as the measure of a dam's global-warming impact. In presenting Belo Monte as an ideal dam from a global-warming perspective, Luis Pinguelli Rosa and coworkers (1996) calculated this ratio as slightly exceeding 10 W/m^2 , considering the originally planned reservoir area of 1225 km^2 (the index would be 25 W/m^2 under the same assumptions considering the currently planned 440 km^2 area).

The regulations of the Kyoto Protocol's Clean Development Mechanism (CDM) currently allow carbon credit for large dams without restriction, but the Executive Board of the CDM, at its meeting in Buenos Aires in December 2004, proposed that these credits be restricted to dams with power densities of at least 10 W/m^2 of reservoir area (UN-FCCC 2004, p. 4), coincidentally the mark achieved by Belo Monte as calculated by Rosa and others (1996). The possibility of claiming carbon credit for Belo Monte

has been raised on several occasions both by Brazilian and World Bank officials (personal observation). A power density as high as 10 W/m^2 for Belo Monte requires that this dam be considered in isolation from the Babaquara (Altamira) Dam that would regulate streamflow at the Belo Monte site by storing water upstream. The two dams together, with $11,000 + 181.3 + 6274 = 17,455 \text{ MW}$ of installed capacity from $440 + 6140 = 6580 \text{ million m}^2$ of reservoir area represents 2.65 W/m^2 of reservoir. This is not much better than the power density of Tucuruí-I (1.86 W/m^2), and far below the 10 W/m^2 limit proposed for Kyoto credit.

In the case of Belo Monte, two reasons make power density highly misleading as a measure of the project's greenhouse impact. First, surface emissions (which are proportional to reservoir area) represent only a part of the global-warming impact of hydroelectric projects; the amounts of methane released from water passing through the turbines (and spillway) depend heavily on the volumes of water that pass through these structures, which (as at Belo Monte) can be large even when reservoir area is small. The second reason is that most of the impact of the overall project is from the upstream dams. In order to fulfill their role in storing and releasing water for use by Belo Monte during the dry season, the upstream dams must be managed with as wide a fluctuation as possible in their water levels. After all, if they were left as "run-of-the-river" dams (*i.e.*, without fluctuations of the reservoir water levels) then the result would be no better than the unregulated river from the point of view of increasing the output of Belo Monte. It is this fluctuation in water level that makes the upstream dams such large potential sources of greenhouse gases, especially Babaquara. The water level in the Babaquara reservoir is expected to vary by 23 m over the course of the year (Brazil, ELETRONORTE nd [C. 1989]). For comparison, the water level in the Itaipú Reservoir on the Brazil/Paraguay border only varies by 30-40 cm. Each time the water level in Babaquara descends to its minimum normal operating level it would expose a vast mudflat of 3580 km^2 (approximately the size of the entire Balbina reservoir!). Soft, easily decomposed vegetation quickly grows in the drawdown zones of reservoirs. When the water level subsequently rises, the biomass decays on the bottom of the reservoir, producing methane.

Reservoirs are thermally stratified, with a boundary (thermocline) typically located at 2-3 m depth. The water temperature abruptly decreases below the thermocline, and water trapped below this layer does not mix with the surface water. This deep water (the hypolimnion) quickly becomes anoxic, and the soft vegetation from the drawdown zone that decomposes under these conditions produces methane (CH_4) rather than carbon dioxide (CO_2). A ton of CH_4 has 21 times more impact on global warming than a ton of CO_2 using the conversion factor (global warming potential, or GWP) adopted by the Kyoto Protocol (Schimel and others 1996), or 23 times more if the most recent value calculated by the Intergovernmental Panel on Climate Change (IPCC) is used (Ramaswamy and others 2001, p. 388). Per metric ton (megagram = Mg) of carbon released in each form, CH_4 has 7.6 times more impact than CO_2 when calculated using the GWP of 21.

The wood in the submerged trees is not believed to be a significant carbon source for methane production because lignified plant tissue (wood) decays extraordinarily slowly under anaerobic conditions. Trees are still usable as timber after several decades if they remain continually submerged, as is shown by the experience at Tucuruí, which was filled in 1984 and 20 years later is still the scene of disputes

between various claimants engaged in exploiting the underwater timber stock. In contrast, soft, green vegetation decomposes quickly, thus releasing its carbon stock in the form of gases, some of which are released to the atmosphere.

The regrowth of vegetation in the reservoir's drawdown zone each year removes carbon dioxide from the atmosphere through photosynthesis, and re-emits the carbon in the form of methane when the vegetation is flooded. The reservoir therefore acts as a virtual methane factory, continually converting CO_2 to CH_4 . The carbon source from the annual flooding of the drawdown zone is permanent, unlike the carbon from the original-forest leaves and leaf litter and labile soil organic carbon. These carbon pools decay over the first few years after the reservoir is filled. Macrophytes (water weeds), another source of easily decomposed biomass, decline to lower levels when the fertility of the water reaches a lower equilibrium after the initial nutrient flush that follows a reservoir filling. Hydroelectric-dam emissions are much higher during the first few years, both from CH_4 generated from underwater decay of soft biomass in the reservoir and from CO_2 from decay of the above-water portions of the original forest trees left standing in the reservoir. Nevertheless, the continual supply of soft biomass from the drawdown zone and from macrophytes guarantees a certain level of permanent emission. The vast drawdown zone of Babaquara assures that this source will be significant.

Features of the Belo Monte and Babaquara Dams

Belo Monte

The design of the Belo Monte reservoir is highly unusual, and greenhouse-gas calculations must be tailored to these features. The reservoir is divided into two independent portions. The "Xingu Channel Reservoir" occupies the course of the Xingu River above the main dam at Sitio Pimentel (Figure 1). The main spillway draws water from this reservoir, as does a small "complementary powerhouse" (181.3 MW installed capacity), which, at periods of peak flow, will make use of some of the water that cannot be used by the main powerhouse. Most of the water will be diverted from the side of the Channel Reservoir through adduction canals to the Canals Reservoir, at the end of which are the intakes for the turbines in the main powerhouse (11,000 MW). The Canals Reservoir also has a small spillway for emergency purposes. Features of the reservoirs are presented in Table 1.

[Figure 1 and Table 1 here]

In order to supply water to the main powerhouse turbines at their capacity of $13,900 \text{ m}^3/\text{second}$, water entering the adduction canals will flow at an average rate of $7.5 \text{ km}/\text{hour}$ in a channel 13 m deep and will take approximately 2.3 hours to make the 17-km trip from the Channel Reservoir to the Canals Reservoir. This will be similar to a river rather than a reservoir. The Canals Reservoir, through which the water will take an average of 1.6 days to pass, is of a form that is perhaps unique in reservoir-construction history. Instead of the usual flooded valley, where water flows through the reservoir following the natural downward-sloping topography of a river and its tributaries, the water in the Canals Reservoir will be flowing across a series of former valleys perpendicular to the normal direction of water flow. The water will pass between five different watersheds as it flows across the courses of the streams that will

have been flooded, passing through shallow bottlenecks as it crosses each of the former interfluvies. Each of these passages, some of which are in channels deepened as part of the construction effort, will offer the opportunity to break any thermocline that may have formed in the deeper valley bottoms in between. It is possible that only relatively well-oxygenated (and low-methane) water from the surface will make the passage through these bottlenecks, leaving relatively permanent pools of methane-rich water at the bottom of each valley. In effect, the 60-km long Canals Reservoir is really a chain of five reservoirs, each with a distinct turnover time, system of associated “dead arms” and potential for stratification. When the water reaches the final stretch before the turbine intakes, it will remain there only for a short while.

Babaquara (Altamira)

In contrast to the small reservoir volume and rapid turnover times of the two reservoirs at Belo Monte proper, the Babaquara reservoir has several features that make it exceptionally noxious as a methane source. One is its huge area – the size of Tucuruí and Balbina put together. Another is the extraordinarily large drawdown area that will be alternately flooded and exposed: 3580 km² (Brazil, ELETRONORTE nd [C. 1989]).

The Babaquara reservoir is divided into two arms, one of which will have a very slow turnover time. The reservoir will flood both the Xingu and Iriri River valleys. Rough measurements of the areas of these sections of the reservoir (from a map in Brazil, ELETRONORTE nd [C. 1988]) indicate approximately 27% of the reservoir area in the Xingu basin below the confluence of the two rivers, another 27% in the Xingu Basin above the point of confluence, and 26% in the Iriri River basin. The average flow (1976-1995) of the Iriri river is 2667 m³/s (Brazil, ANEEL 2001), while the flow at the Babaquara Dam site (*i.e.*, below the confluence) is 8041 m³/second (Maceira and Damázio nd). Assuming that the portion of the reservoir below the confluence (the portion nearest the dam) is three times as deep, on average, as the other two segments, then the residence time in the Babaquara reservoir for water descending the Xingu is 164 days, while for water descending the Iriri it is 293 days. While the residence time is very long in either case, giving ample time to accumulate a large load of methane, the time for the Iriri portion almost reaches that of the notorious Balbina Dam’s 355-day residence time (Fearnside 1989)! The tremendous difference between Babaquara and Belo Monte, with vertical fluctuations in water levels varying from zero in the Belo Monte Channel Reservoir to 23 m at Babaquara, indicate that an explicit model of carbon stocks and degradation is needed, rather than a simple extrapolation of measurements of CH₄ concentrations and emissions at other dams. The model developed for this purpose is described in the following sections.

Carbon sources and Greenhouse-Gas Release Pathways

Methane

Methane produced by underwater decay can be released in various ways. One is by bubbling and diffusion through the reservoir surface. Bubbling, which allows CH₄ to pass through the barrier of the thermocline, is highly dependent on the depth of the water at each point in the reservoir, bubbling emissions being much greater at shallower depths. Diffusion is important in the first year, but not thereafter; this is because

bacterial populations in the surface water (epilimnion) then increase, with the result that any methane diffusing through this layer is oxidized to CO₂ before it reaches the surface (Dumestre and others 1999, Galy-Lacaux and others 1997). The surface emissions are also higher in the first years after filling because the leaves and leaf litter from the original forest and the labile portion of the soil carbon is being released from the bottom of the reservoir as methane. These initial carbon stocks will decline as they are progressively exhausted and, in later years, carbon will only be available from renewable carbon sources such as macrophytes and the drawdown zone regrowth (as well as soil carbon entering the reservoir from upstream erosion).

Studies to quantify the relative role of different carbon sources are lacking. At the Petit Saut reservoir in French Guiana, Galy-Lacaux and others (1999) believe that soil carbon is the principal source in the first years. The stock of labile soil carbon is large relative to the other stocks of easily degraded carbon. The present calculation uses the labile (hydrosoluble) soil carbon stock of 54 Mg C/ha measured in the top 60 cm of a typical Amazonian Ultisol (Trumbore and others 1990, p. 411). Assumptions regarding the rate of decay of the stocks produce a theoretical total for the carbon released into the water as CH₄. Considering the dilution effect of inflows to the reservoir, the amount of anaerobically decomposed carbon per billion cubic meters of water can be calculated. This calculated amount has been derived for two existing tropical-forest reservoirs (Petit Saut and Tucuruí) and related to the CH₄ concentration in the water at a standardized depth (30 m) in the same reservoirs.

The amount of carbon decayed anaerobically is the sum of the portions decayed of original leaves and leaf litter, labile soil carbon, unbeached macrophytes and flooded drawdown vegetation. The amount of water is the reservoir volume at the end of the month plus the inflows during the month and the previous month. The relation of the amount of carbon decayed anaerobically (calculated according to the assumptions given above) to CH₄ concentration at 30 m depth is shown in Figure 2; concentration data are from Petit Saut (Galy-Lacaux and others 1999), with the exception of the point at the far left with 6 mg CH₄/liter at 30-m depth, which is from Tucuruí (J. G. Tundisi, cited by Rosa and others 1997, p. 43). The range of values for the amount of carbon decayed anaerobically is divided into three segments for calculation of CH₄ concentration at 30 m depth (equations 1-3).

[Figure 2 here]

For anaerobic decay ≤ 684.4 Mg C/billion m³ of water:

$$Y = 0.00877 X \quad (\text{eq. 1})$$

For anaerobic decay 684.5 – 15,000 Mg C/billion m³ of water:

$$Y = 0.000978 X + 6 \quad (\text{eq. 2})$$

For anaerobic decay $> 15,000$ Mg C/billion m³ of water:

$$Y = 20 \quad (\text{eq. 3})$$

Where: X = anaerobic decay (Mg C/billion m³ of water)

$Y = \text{CH}_4 \text{ concentration at 30 m depth (mg/liter)}$

The ratio of the methane concentration at different depths to the concentration at 30 meters depends on the age of the reservoir, since the relationship changes over time as the bacterial populations in the surface waters become more capable of degrading methane to carbon dioxide. Data from the Samuel reservoir when five months old (J.G. Tundisi, cited by Rosa and others 1997, p. 43) are used to represent reservoirs up to 12 months after filling; data from Petit Saut (Galy-Lacaux and others 1999) are used to represent reservoirs from the 13th to the 36th month, and data from Tucuruí collected 44 months after filling (J.G. Tundisi, cited by Rosa and others 1997, p. 43) are used to represent reservoirs after month 36. The ratios are calculated using the equations in Table 2.

[Table 2 here]

Bubbling and diffusion emissions can be related to the concentration at the standardized depth of 30 m. Table 3 presents equations for these emissions for areas with different water depths. These relationships have been derived from the measurements at Petit Saut (Galy-Lacaux and others 1999). The predicted CH_4 concentration at 30 m depth is closely related to the observed bubbling emissions in each depth range in the Petit Saut data (0-3 m, 4-6 m and 7-8 m) (Figure 3a, b and c). Diffusion emissions at Petit Saut, independent of depth, are also closely related to the predicted CH_4 concentration at 30 m (Figure 3d).

[Table 3 and Figure 3 here]

Using the ratios derived above, the CH_4 concentration in Babaquara and the two Belo Monte reservoirs can be calculated. The calibration of calculated carbon release through anaerobic decay to the existing data on CH_4 concentration in similar reservoirs is important in reducing possible bias from the assumptions regarding the size and decay rates of the various underwater carbon stocks. The water entering a reservoir from streams and from normal river flow, such as that entering Babaquara, contains virtually no CH_4 , as shown by measurements at Petit Saut (Galy-Lacaux and others 1997). In the case of Belo Monte, however, water entering directly from Babaquara will contain appreciable amounts of CH_4 .

The water management at Babaquara is assumed to follow a logic based on supplying the maximum amount of water annually to Belo Monte, within the limitations posed by the seasonal cycle of river flows, the maximum that can be used by the turbines at Babaquara, and the live storage volume of the reservoir. This results in the expected annual rise and fall in the water level. For each month over a 50-year period a calculation is made of the area of drawdown zone that has remained exposed for one month, two months, and so forth up to one year, and a separate category is maintained for area of drawdown exposed for over one year. The area that is submerged in each age class is calculated for each month. This allows a calculation of the amount of soft biomass that is flooded, based on assumptions regarding the growth rate of the vegetation in the drawdown zone. The category for vegetation over one year old contains less soft biomass, as growth after the first year would be largely allocated to producing wood rather than more soft material (the leaf biomass of the forest is used for this category).

Macrophytes are an important source of soft, easily decomposed biomass. The populations of these aquatic plants explode to cover a substantial part of a new reservoir, as occurred at Brokopondo in Surinam (Paiva 1977), Curuá-Una in Pará (Junk and others 1981), Tucuruí in Pará (de Lima 2002), Balbina in Amazonas (Walker and others 1999), and Samuel in Rondônia (Fearnside in press). LANDSAT satellite imagery indicates that at Tucuruí macrophytes covered 40% of the reservoir surface two years after filling, subsequently declining to 10% a decade later (de Lima and others 2002). Based on monitoring at Samuel and Tucuruí, Ivan Tavares de Lima (2002) developed an equation (eq. 4) to describe the path of macrophyte cover, which is used in the present analysis:

$$Y = 0.2 X^{-0.5} \quad (\text{eq. 4})$$

where:

X = Years since flooding

Y = The fraction of the reservoir covered by macrophytes.

Macrophytes die at a given rate in the reservoir and the dead biomass sinks to the bottom. In *várzea* (floodplain) lakes, macrophyte death results in a turnover of the biomass 2-3 times per year (Melack and Forsberg 2001, p. 248); the midpoint of this range (4.8 months) implies that 14.4% of the macrophyte biomass dies each month. This rate has been adopted for macrophyte mortality in the reservoirs. In addition to this mortality, a part of the macrophyte biomass is beached when the water level falls. Because the prevailing winds (which blow from east to west) push the floating macrophytes against one shore, a part of the carpet of floating of plants is necessarily positioned where it will be beached whenever the water level descends. The quantities involved are impressive, as is evident at Tucuruí (see Fearnside 2001). Because macrophytes concentrate along only one shore of the reservoir, only half of the drawdown zone is considered in computing areas of beached macrophytes. When beached, the macrophytes die and decay aerobically. However, if the water level rises again before the decay process is complete, the remaining carbon stock in beached macrophytes is added to the pool of underwater carbon that can produce methane. Here it is assumed that, if an area is exposed for only one month, then half of the beached macrophytes will still be present when these areas are re-flooded.

The macrophyte cover in Amazonian reservoirs undergoes a regular sequence of species succession, beginning with *Eichhornia* and ending with *Salvinia*, as occurred at Curuá-Una (Vieira 1982) and Balbina (Walker and others 1999). *Eichhornia* and other species that predominate in the early years have significantly greater biomass per hectare than *Salvinia*. At Balbina the replacement of higher-biomass macrophytes by *Salvinia* occurred between the seventh and eighth year after filling (Walker and others 1999, p. 252). The shift to *Salvinia* is assumed to occur seven years after reservoir filling in the present calculations for the Xingu dams. Floating macrophytes such as *Eichhornia* and *Salvinia* are most common in reservoirs, but some rooted species also occur.

For the first six years, the biomass of macrophytes is assumed to be 11.1 Mg/ha dry weight, based on an *Eichhornia* mat measured at Lago Mirití, a *várzea* lake near Manacapuru, Amazonas (P. M. Fearnside, unpublished data). For comparison, *Oryza*

species (a rooted grass) in *várzea* lakes, had 9-10 Mg/ha of dry weight, while *Paspalum* had 10-20 Mg/ha (T.R. Fisher, D. Engle and R. Doyle, unpublished data cited by Melack and Forsberg 2001, p. 248). In another measurement in *várzea* lakes (where nutrient availability is greater than will be the case in the Xingu dams), nine measurements of rooted macrophytes in the *várzea* after approximately three months of growth averaged 5.7 Mg/ha (SD=1.7, range=3.2-8.7) (Junk and Piedade 1997, p. 170). The value of 11.1 Mg/ha in the Xingu dams is similar to values for floating and submerged macrophyte biomass in other parts of the world. For example, the submerged macrophyte load in Lake Biwa, Japan has 7-10 Mg/ha of dry biomass (Ikusima 1980, p. 856).

After the transition to *Salvina* takes place the biomass per hectare of macrophytes is lower. The biomass value used in the calculation is 1.5 Mg/ha dry weight, which is the biomass of mats of *Salvinia auriculata* in *várzea* lakes (Junk and Piedade 1997, p. 169).

The methane in the water that is trapped below the thermocline will be exported from the reservoirs in the water drawn by the turbines and the spillway. This is a feature of hydroelectric dams that is completely different from natural water bodies such as *várzea* lakes, which are globally significant sources of CH₄ from surface emissions alone. Opening the intakes for the turbines and spillway is like pulling the plug in a bathtub—the water is drawn from the bottom, or at least from the bottom portion (hypolimnion) of the reservoir. Below the thermocline the concentration of CH₄ increases steadily as one descends through the water column. An important observation from Petit Saut is that, within a reservoir, the CH₄ concentration at a given point in time is approximately constant at any given depth below the surface—regardless of the depth to the bottom at the location in question (Galy-Lacaux and others 1997). The present analysis tracks the depth below the water surface of the spillway and turbine intakes in order to calculate the corresponding CH₄ concentration in the water released through these structures.

As one descends through the water column, the pressure increases and the temperature decreases. Both effects act to increase the concentration of CH₄ at greater depths. By Henry's Law, the solubility of a gas is directly proportional to the pressure, while Le Chatelier's Principle holds that gas solubility is inversely proportional to temperature. While both effects are important, the effect of pressure predominates (Fearnside 2004). The pressure is almost five atmospheres at the 48-m turbine intake depth at the normal operating level in Babaquara. When the water emerges from the turbines, the pressure instantly drops to one atmosphere. Dissolved gases are released when the pressure drops, just as bubbles of CO₂ emerge immediately when one opens a bottle of Coca Cola. The pressure drop when a bottle of Coca Cola is opened is much less than the pressure drop when water emerges from the turbines of a hydroelectric dam, thus making the degassing even more immediate. The ease with which each gas comes out of solution is determined by the Henry's Law constant of the gas. This constant is higher for CH₄ than for CO₂, so the methane would also be released more readily than the bubbles from a bottle of Coca Cola for this reason. At Petit Saut, for example, the water entering the turbines in 1995 had a ratio of dissolved CO₂ to CH₄ of 9:1, but in the plume immediately below the dam the ratio was 1:1, meaning that proportionally much more of the dissolved methane had been released (Galy-Lacaux and others 1997).

The fraction of the dissolved CH₄ that is released in water passing through the spillway and turbines will depend on the configuration of each. In the case of the spillway at Babaquara, the fall of 48 m after emerging from the floodgates (Table 1) should guarantee virtually complete release. In the case of the turbines, however, some of the CH₄ content will probably be passed on to the Belo Monte reservoir immediately below Babaquara. The Belo Monte reservoir is planned to back up against the bottom of the Babaquara Dam, such that the water emerging from the Babaquara turbines will be directly injected into the Belo Monte reservoir itself, rather than flowing in a stretch of normal river before entering the reservoir. Because the water drawn from deep in the water column of the Babaquara reservoir will be at low temperature, it will probably sink immediately into the hypolimnion once it enters the Belo Monte reservoir. Its CH₄ content would therefore be partially preserved—and would be subject to release when the water later emerges from Belo Monte's turbines.

Carbon dioxide

Unlike methane, carbon dioxide is removed from the atmosphere through photosynthesis when plants grow. The CO₂ released from decay of soft biomass that has grown in the reservoir and its drawdown zone therefore cannot be counted as a global-warming impact, as this is merely being cycled repeatedly between the biomass and the atmosphere. The biomass in the forest trees that were killed when the reservoir was created is a different matter, and the CO₂ it releases constitutes a net impact on global warming. Only the above-water portion of this biomass decays at an appreciable rate.

Above-water wood biomass is modeled in some detail, based on what is known from the experience at Balbina (which filled over the 1987-1989 period). Trees break off just above the high-water mark. By eight years after flooding, approximately 50% of the trees ≥ 25 cm in diameter and 90% of the trees < 25 cm in diameter had broken (Walker and others 1999). In addition, branches continually fall from the standing trees. Approximately 40% of *terra firme* (upland) trees float in water (see Fearnside 1997a). The trees that sink stay where they fall, either in the permanently flooded zone or in the shallower areas that are periodically exposed as the drawdown zone. Those that float are pushed by wind and waves to the shore and will be exposed to aerobic decay in the drawdown zone when the water level descends. The stocks and decay rates in each category are calculated. Aerobic decay contributes to the CO₂ emissions from above-water biomass. Parameters for the dynamics and aerobic decay of above-water biomass are given in Table 4.

[Table 4 here]

The above-water biomass emissions considered here are conservative for two reasons. One is that they are based on average water flows for each month and on the assumption that water management respects the limit of the minimum normal water level foreseen for the reservoir. No consideration is given to the possibility that the level might be drawn down below this level in extremely dry years, as in El Niño events. The second conservative assumption is that biomass in the drawdown zone never burns. Burning is only an occasional event, but it affects significant amounts of biomass when it does occur. During the 1997-1998 El Niño drought, the reservoirs of

both the Balbina and Samuel dams were drawn down to elevations much lower than their official “minimum” levels, and large areas of the expanded drawdown zones burned. Although it is likely that such emissions will sometimes occur at Babaquara, they are not considered in this analysis.

Another source of emissions is from trees near the edge of the reservoir that are killed when the water table rises and reaches their roots. At Balbina, a band of dead trees is evident all around the edge of the reservoir (Walker and others 1999). Because the format of the shoreline is exceedingly tortuous and includes the edges of many islands created by the reservoir, this band of forest dieback encompasses a significant area. The dead trees decay, releasing CO₂, and over a period of decades secondary forest develops (with an attendant absorption of carbon). The present analysis assumes that mortality is 90% within 50 m of the reservoir edge and 70% if 50-100 m from the edge. Decay follows the same course as in areas felled for agriculture, while secondary forests are presumed to grow at the same rate as those in shifting-cultivation fallows (Fearnside 2000).

Pre-Dam Ecosystem Emissions

The emissions of ecosystems present before the dams were built must be deducted from the dams’ emissions in order to have a fair evaluation of the net impact of the hydroelectric development. The idea that the forests flooded by the reservoir have large natural emissions of greenhouse gases has been a major component of the attack that the hydropower industry has mounted against studies indicating high emissions from hydroelectric dams. When early studies indicated that the Balbina Dam emitted more than would be released by producing the same amount of electricity from fossil fuels (Fearnside 1995a), the US National Hydropower Association (USNHA) reacted with the statement:

“It’s baloney and it’s much overblown ... Methane is produced quite substantially in the rain forest and no one suggests cutting down the rain forest.”

This statement by Karolyn Wolf (spokesperson for USNHA) illustrates the vehemence with which this subject has been resisted (see IRN 2002). Hydro-Québec even went so far as to assert that large emissions from floodplain ecosystems in the areas flooded by hydroelectric dams could make the net impact of these projects a “zero-sum issue” (Gagnon 2002). Unfortunately, a closer examination of these arguments points instead to a major net emission from hydroelectric dams. Babaquara illustrates this well, and it is worth examining this case in some detail.

The areas of flooded and unflooded ecosystems are given in Table 5. The seasonally flooded forest types are taken as the area flooded, but this may be an overestimate since radar imagery from the Japanese Earth Resource Satellite (JERS) indicates that virtually none of the reservoir area has flooding below the forest canopy (see Melack and Hess 2004). It should be noted, however, that temporary oxbow lakes along the Xingu and Iriri rivers do exist: maps analyzed by de Miranda and others (1988, p. 88) indicate from 28 to 52 such lakes in the area to be flooded by Babaquara, depending on the map used in the analysis.

[Table 5 here]

The parameters for methane emissions by unflooded forest are given in Table 6. These indicate a minimal effect on methane, with a small sink coming in the soil being lost to flooding. Nitrous oxide (N₂O) emissions from unflooded forest soil are small: 0.0087 Mg gas/ha/year (Verchot and others 1999, p. 37), or 0.74 Mg/ha/year of CO₂-equivalent carbon considering the global warming potential of 310 (Schimel and others 1996, p. 121). Nitrous oxide calculations for unflooded forest and flooded areas are given in Table 7. The calculations include the effect of temporary ponding on *terra firme* during periodic heavy rainfall events (Table 7).

[Tables 6 and 7 here]

For flooded areas, the assumption is made that each flooded point is submerged for an average of two months per year. Of course some parts of the area would be submerged for longer and some for shorter times, depending on the altitude of each point. The value used for emissions per hectare (103.8 mg CH₄/m²/day, SD=74.1, range=7-230) is the mean of five studies in white-water *várzea* forest reviewed by Wassmann and Martius (1997). A similar value of 112 mg CH₄/m²/day while flooded (n=68, sd=261) was found in blackwater flooded forests (*igapós*) along the Jaú River, a tributary to the Rio Negro Basin. In the *igapó* forests in the Jaú basin studied by Rosenqvist and others (2002, p. 1323) the rate of methane emission from the flooded areas is much higher during the short period when the water level is falling than it is during the remainder of the time that the area is under water. This would tend to make the annual emission somewhat independent of the time period that the areas are flooded, and makes the result relatively robust for extrapolation to other river basins in Amazonia if expressed in terms of emission per flooding cycle (rather than per day flooded). Assuming the same emission rates as those measured in the white-water *várzea* studies (the Xingu is considered a clear-water river, more similar to white water than black), the annual emission would be equivalent to only 0.043 million tons of CO₂-equivalent carbon at Babaquara on a per-day basis, or 0.248 million tons of CO₂-equivalent carbon if this result is multiplied by 3 to approximate the effect of the shorter (2-month versus 6-month) flooded season. The resulting adjustments from pre-dam ecosystems are very minor, as will be shown later when net emissions are calculated for the two dams.

Construction Emissions

Dams obviously require much more materials, such as steel and cement, than do equivalent fossil-fuel powered facilities such as the gas-fired powerplants that are now being built in São Paulo and other cities in central-south Brazil. The quantities of steel to be used in the construction of Belo Monte are calculated based on the weights of the items listed in the feasibility study (Brazil, ELETRONORTE 2002). For Babaquara, the amount of steel to be used in electro-mechanical equipment is assumed to be proportional to installed capacity, while the amount of steel in concrete reinforcing rods is assumed to be proportional to the volume of concrete (from da Cruz 1996, p. 18). The quantities at Babaquara are calculated in proportion to the amounts used at Belo Monte. Estimated steel use at Belo Monte totals 323,333 Mg, while that at Babaquara totals 303,146 Mg.

The amount of cement to be used at Belo Monte is estimated at 848,666 Mg, based on the total of the items listed in the feasibility study (Brazil, ELETRONORTE 2002). For Babaquara, cement use is estimated at 1,217,250 Mg based on the volume of concrete (from da Cruz 1996, p. 18) and an assumed average cement content of 225 kg/m³ of concrete (Dones and Gantner 1996). Belo Monte is unusually sparing in cement use because the site allows the main dam (Sitio Pimentel) to be built at a location that is higher in elevation than the main powerhouse (Sitio Belo Monte). The main dam has a maximum height of only 35 m (Brazil, ELETRONORTE 2002, Tomo I, p. 6-33), while the main powerhouse benefits from a reference fall of 87.5 m (Brazil, ELETRONORTE 2002, Tomo I, p. 3-52). Most hydroelectric projects, such as Babaquara or Tucuruí, have the powerhouse located at the foot of the dam itself, and so only generate power from a fall that corresponds to the height of the dam minus a small allowance for freeboard at the top. Tucuruí, which is so far the “champion” of all Brazilian public works projects in terms of cement use, used three times as much cement as that required for Belo Monte (Pinto 2002, p. 39). Babaquara uses 2.6 times more cement per MW of installed capacity than Belo Monte.

The amount of diesel used for Belo Monte is expected to be 400×10^3 Mg (Brazil, ELETRONORTE 2002, Tomo II, p. 8-145). This includes an adjustment of the units (as reported in the feasibility study) to bring the values within the ballpark of fuel use at other dams (*e.g.*, Dones and Gantner 1996 calculated a mean use of 12 kg diesel/TJ for dams in Switzerland); the feasibility study contains a variety of internal inconsistencies in units that presumably stem from typographical errors. Belo Monte has an unusually large amount of excavation because of the need to dig the adduction canal connecting the Channel and Canals Reservoirs, and the various smaller excavation projects at the bottlenecks within the Canals Reservoir. The expected amount of excavation for these canals increased substantially between the 1989 and 2002 versions of the feasibility study because errors were discovered in the topographic mapping of the area (Brazil, ELETRONORTE 2002, Tomo I, p. 8-22). For Babaquara, diesel use is assumed to be proportional to the amount of planned excavation at that dam (da Cruz 1996, p. 18).

The estimates of materials for construction of dams and transmission lines are given in Table 8. The resulting totals (0.98 million Mg C for Belo Monte and 0.78 million Mg C for Babaquara) are minuscule compared to the later emissions from the reservoirs. The construction emissions from equivalent gas-fired plants have not been deducted from these construction-emission totals. Construction emission from natural gas facilities is minimal: life-cycle analysis of combined-cycle gas plants in Manitoba, Canada indicates construction emissions of only 0.18 Mg CO₂ equivalent/GWh (McCulloch and Vadgama, 2003, p. 11).

[Table 8 here]

Calculated Emissions from Belo Monte and Babaquara

Calculation of greenhouse-gas emissions requires a realistic scenario for the timing of filling and turbine installation events at Belo Monte and Babaquara, and for reservoir management policies at the two dams. Here it is assumed that Babaquara is filled seven years after Belo Monte (*i.e.*, that Belo Monte operates using unregulated water flow before this time). This schedule corresponds to the less-optimistic scenario

in the original plan (see Sev 1990). The turbines in both dams are installed at a rate of one every three months, the (perhaps optimistic) pace foreseen in the feasibility study (Brazil, ELETRONORTE 2002, Tomo II, p. 8-171).

The present calculation follows the plans for reservoir filling as given in the feasibility study. The Canals Reservoir will first be filled to a level 91 m above mean sea level. This will be done after the first flood passes through the spillway (Brazil, ELETRONORTE 2002, p. 8-171). This is assumed to occur in the month of July. The complementary powerhouse will then be used at this reduced reservoir level for one year before the main powerhouse is ready for use, as planned in the ELETROBRS Decennial Plan (Brazil, MME-CCPESE 2002). The Decennial Plan’s reference scenario calls for the complementary powerhouse to begin operation in February 2011 and the main powerhouse in March 2012.

The results of a 50-year calculation of the sources of soft, easily degraded carbon are shown for each reservoir in Figure 4. It is evident that all sources are much higher in the first years than they are in later years. The stocks of labile soil carbon, above-water wood biomass and dead trees along the shoreline dwindle, consequently reducing the emissions from these sources. Macrophytes decline, but do not disappear, thus providing a long-term source that, in the later years, is of greater relative importance, albeit smaller in absolute terms. Regrowth in the drawdown zone represents a stable long-term source of degradable carbon, which increases in relative importance as the other sources decline.

[Figure 4 here]

The calculated methane concentrations at a standardized depth of 30 m are shown for each reservoir in Figure 5. These calculated concentrations follow the general trend of seasonal oscillation and asymptotic decline observed in measured values at Petit Saut (Galy-Lacaux and others 1999, p. 508). The oscillations are very large at Babaquara after the carbon sources other than drawdown vegetation have declined in importance (Figure 5a). The large peaks in methane concentration at Babaquara are maintained after the concentrations during the remainder of each year have declined; the high peaks are maintained because the carbon comes from flooding of drawdown vegetation when the water rises. The concentration peaks result in substantial emissions because these periods correspond to periods of high turbine flow in order to maximize power output.

[Figure 5 here]

The emissions through each pathway for the Belo Monte/Babaquara complex as a whole are shown in Figure 6. Above-water biomass and shoreline dieback decline to insignificant levels over the 50-year time span, but the large magnitude of emissions from above-water biomass in the early years gives this source a significant place in the overall 50-year average. Fifty years is the time period generally adopted by the hydroelectric industry in discussing the “useful life” of dams, and calculations, both financial and environmental, are often made on this time span, as in the regulations applying to feasibility studies for dams in Brazil (Brazil, ELETROBRS and DNAEE 1997). The existing Amazonian dams, particularly Tucuru, Balbina and Samuel, were relatively young in 1990—the worldwide standard year for the baseline greenhouse gas

inventories mandated by the United Nations Framework Convention on Climate Change and the year used for several previous calculations of greenhouse-gas emissions (Fearnside 1995a, 1997b, 2002a, in press). The emissions in 1990 were therefore quite high, and the hydroelectric industry has frequently objected that these estimates give an unfairly negative picture of the role of hydropower in global warming (*e.g.*, IHA and [C. 2002]). The current calculation shows that even over a 50-year time horizon the global-warming impact of a dam like Babaquara is significant.

[Figure 6 here]

Long-term averages of net emissions of greenhouse gases are presented in Table 9 for different time horizons. Emissions are separated into those considered under hydroelectric dams in the national inventories that are being prepared by each country under the Climate Convention (UN-FCCC) and other fluxes that are also part of the net impact and benefit of the dam, including avoided emissions. The total impact of the dams averages 11.2 million of CO₂-equivalent carbon. for the 1-10 year period, decreasing to 6.1 million Mg for the 1-20 year period and -1.4 million Mg for the 1-50 year period.

[Table 9 here]

Key Uncertainties

A calculation such as the present one for the Belo Monte/Babaquara complex involves a great deal of uncertainty. It must be done nevertheless, and the best information available must be used for each of the parameters required by the model. As research in this area proceeds, better estimates for these parameters will become available and the model can quickly interpret these findings in terms of their effect on greenhouse-gas emissions.

Although a full set of sensitivity tests has not yet been conducted, the behavior of the model provides a number of indications of which parameters are most important. Sensitivity tests for selected input parameters are presented in Table 10, showing the effect of a 10% increase in each input parameter. Effects are symmetrical for a 10% decrease in each parameter (not shown in the table). Effects are presented in terms of the percentage change in the total impact of the dams as annual averages for the 1-10 year, 1-20 year and 1-50 year periods, that is, as a percentage departure from the reference scenario values for these averages as presented in Table 9. For all three periods, the variables to which the total impact is most sensitive are the biomass of the original forest and the percentages of the exported methane that is emitted at the turbines and at the spillways.

[Table 10 here]

In the early years after reservoir filling, emissions are dominated by CO₂ released by above-water biomass decay. These emissions, while subject to uncertainty, are founded on the best available data from decomposition in deforested areas. While measurements specific to reservoir trees would be valuable, a radical change in the result is not expected. The mortality assumptions for the forest at different distances

from shoreline are no more than guesses, but in this case the amount of carbon involved is insufficient to make any significant difference in the overall result.

The initial years also include a substantial emission from release of methane in water passing through the turbines. The percentage of the dissolved methane released at this point is a critical parameter. High- and low-emissions scenarios were run, using the values derived from available measurements at Petit Saut (Galy-Lacaux and others 1997, 1999). Because of differences between Petit Saut and the Brazilian dams, the range used is very wide (21-89.9%) (See discussion in Fearnside 2002a). The emissions estimates presented here are the midpoints of the results produced with the high- and low-end values for the percentage emitted at the turbines. This middle value is believed to be conservative.

It should be noted that, when both Belo Monte and Babaquara are in operation there is a certain compensation between the two dams that reduces the overall effect of uncertainty regarding the percentage of dissolved methane released at the turbines. When a low estimate for this parameter is used, the emission at Babaquara is reduced but the unreleased CH₄ is passed on to Belo Monte, where emissions through other pathways (surface emissions and emissions in the adduction canal and bottlenecks) consequently increase.

The carbon sources for CH₄ emissions in the early years are dominated by release of labile soil carbon stocks (Figure 4). While measurements of this release are lacking in any reservoir, the scaling of the 30-m depth CH₄ concentration to observed values in this range of early years, especially at Petit Saut, results in a realistic trajectory of CH₄ concentrations and emissions from this source.

Most important are the uncertainties regarding CH₄ emission after the initial peak has passed. Much less data from older Amazonian reservoirs are available to calibrate this part of the analysis. The decline in macrophyte areas reduces the importance of the uncertainty regarding this source for the long-term emissions. What predominates for the complex as a whole is the biomass from the drawdown zone in Babaquara. This results in large seasonal peaks in the CH₄ concentration in the Babaquara reservoir (Figure 5a), some of which is passed on to the two Belo Monte reservoirs (Figures 5b and 5c). The growth rate of the drawdown vegetation is therefore critical, and no actual measurements of this exist in Amazonian reservoir drawdown zones. The assumption made that this growth occurs linearly, accumulating 10 Mg dry matter in one year; the value used for the carbon content of this and other forms of soft biomass is 45%. The assumed growth rate is extremely conservative, when compared to measured growth rates of annual herbaceous plants for the three-month period of exposure in *várzea* floodplain areas along the Amazon River near Manaus: in 9 measurements by Junk and Piedade 1997, p. 170) these plants accumulated an average of 5.67 Mg dry weight/ha (SD=1.74, range 3.4-8.7). The proportional value for one year of linear growth would be 22.7 Mg/ha, or more than double the value assumed for the Babaquara drawdown zone. A measurement of the above-ground biomass of grasses up to 1.6 months after *várzea* is exposed at Lago Mirití indicates a dry-matter accumulation rate equivalent to 15.2 Mg/ha/year (P. M. Fearnside, unpublished data). The soil fertility of *várzea* sedimentation zones is greater than what would apply to a reservoir drawdown zone, but an assumption on the order of half to two-thirds the

várzea growth rate seems safe. Nevertheless, this is a major point of uncertainty in the calculation.

Decomposition rates are also important, and measurements under anaerobic conditions in reservoirs are unavailable. Decomposition of *várzea* herbaceous vegetation is believed to provide an adequate parallel. In measurements under flooded conditions in white-water *várzea*, decay of three species (Furch and Junk 1997, p. 192; Junk and Furch 1991) and an experiment in a 700-liter tank with a fourth species (Furch and Junk 1992, 1997, p. 195) indicated the fraction of dry weight lost after one month of submergence averaged 0.66 (SD= 0.19 range=0.425-0.9). The low value in the range was from the species measured in the tank experiment, where the water was known to be anoxic after only about one day; if the measurements under natural conditions included any aerobic decay, the average rate for truly anoxic conditions might be somewhat lower than the four-species mean used here.

Decay rates for aerobic decomposition of beached macrophytes determine how much of this biomass is still present if the water level rises again before decay is complete. A measurement of dead *Eicchornia* at Lago Miriti up to 1.6 months after beaching indicates a loss of 31.4% of the dry weight per month (P. M. Fearnside, unpublished data). Sample size is minimal (three 1-m² plots).

Water management at Babaquara is also important in determining the amount of emission from the drawdown zone. The longer the reservoir is maintained at a low water level, the more vegetation grows in the drawdown zone. The subsequent release of CH₄ generated by flooding the drawdown zone more than compensates for the effect in the opposite direction from lower water levels reducing the depth to the Babaquara turbine intake and therefore the CH₄ concentration in the water passing through the turbine while the water level is low. The assumptions for water use employed in the calculation result in three months of low water levels, four months of high levels and five months of intermediate levels.

The magnitude of the high seasonal peaks of CH₄ depend on the relationship between the amount of degradable carbon and the CH₄ stock (and concentration) when these variables were at high levels in the early years at Petit Saut (*i.e.*, data from Galy-Lacaux and others 1997, 1999). The nature of the carbon source at Petit Saut during this time was different (believed to be primarily soil carbon). The true amount of carbon degraded anaerobically at Petit Saut during this time is unknown, and the scaling that lends confidence to the results for the initial years after reservoir filling, when the carbon sources were of the same type, does not lend so much confidence to these results in later years. Quantifying the relationship between the amount of decay of soft biomass (such as macrophytes and especially drawdown-zone vegetation) and CH₄ production should be a top research priority. However, the general result, namely that drawdown vegetation produces a large and renewable pulse of dissolved CH₄ in reservoirs, is not in doubt. A case in point is the experience at the Três Marias Dam in the state of Minas Gerais, where a 9-m vertical fluctuation in water level results in exposure and periodic flooding of a large drawdown zone, with a subsequent large peak of surface emissions of methane (Bodhan Matvienko, personal communication 2000). Even at the very advanced age of 36 years, the Três Marias reservoir emits methane through bubbling in amounts that greatly exceed the surface emissions of all other

Brazilian reservoirs that have been studied, including Tucuruí, Samuel and Balbina (Rosa and others 2002, p. 72).

An additional source of uncertainty is the fate of dissolved CH_4 as water passes through Belo Monte's 17-km adduction canal and through the four sets of bottlenecks separating the flooded stream basins that make up the Canals Reservoir. Part of the methane is emitted, part is oxidized, and the remainder is passed on to the Canals Reservoir. The parameters used for this are based on the assumption that the canal (width at the surface approximately 526 m and the flow at full capacity 13,900 m^3/second) is similar to the stretch of French Guiana's Sinnamary River below the Petit Saut Dam (where the average river width is 200 m and the mean flow is only 267 m^3/second). Galy-Lacaux and others (1997) estimated methane concentrations and fluxes along the 40 km stretch of river below the Petit Saut Dam, from which they calculated the amounts emitted and oxidized in the river. Their results indicate that, for the dissolved CH_4 entering the river from the dam, 18.7% is released and 81.3% is oxidized (mean of measurements at three dates, with percentage released ranging from 14 to 24%). Virtually all of the release and oxidation occurs within in the first 30 kilometers. In the Sinnamary River, after a 4-km initial stretch where a mixing process occurs, both CH_4 concentration and surface flux decline linearly to zero at 30 km below the dam (*i.e.*, over a 26-km stretch of river). Considering the stock at each point along the river, one can calculate that in the first 17 km of river, 15.3% of the CH_4 is released, and 66.5% is oxidized. In the calculation for Belo Monte these percentages were assumed to apply, and the remaining methane is passed on to the Canals Reservoir.

Estimates for emission at the bottlenecks were derived from information on their length and the percentages of emission and oxidation that occurred over a stretch of river of the same length below the Petit Saut Dam. Based on a map of the reservoir (Brazil, ELETRONORTE nd [C. 2002]b), the first and second compartments are connected by three passages averaging 1.6 km in length, the second and third compartments are connected by two passages averaging 1.7 km in length, the third and fourth compartments are connected by two passages averaging 1.3 km in length, and the fourth and fifth compartments are connected by a wide passage that (although undoubtedly shallow at the interbasin divide) can be considered as a passage of 0 km in length. The percentages of dissolved methane released and oxidized at these bottlenecks is assumed to be proportional to the percentage of release and oxidation that occurred over this same length of river below the Petit Saut Dam (based on data from Galy-Lacaux and others 1997). Uncertainty in this case is much higher than in the case of the values for these percentages calculated for the adduction canal because the short bottlenecks are within the initial stretch of river where a mixing process was in progress. The percentages used (which are all very low) also assume that the process stops at the end of the bottleneck, rather than continuing for some distance into the next compartment of the reservoir. The net result is that the bottlenecks, taken together, only emit 2.1% of the methane, while 9.2% is oxidized and 88.7% is transmitted on to the end of the reservoir.

As at the Babaquara turbines, there is some compensation in the system for uncertainty in the percentages released at the adduction canal and at the bottlenecks. If the emissions at the adduction canal and/or at the bottlenecks are overestimated, then the emission at the Belo Monte main powerhouse turbines will be underestimated. Note that this only applies to the values for the percentage emitted, not to the values used for

the percentage of oxidation in these canals: any over- or underestimate in the percentage oxidized would not be compensated by any change in the opposite direction in turbine emissions.

In summary, multiple uncertainties exist in the current calculation. Future research, especially if targeted to the parameters to which the model indicates the system is most sensitive, will help reduce these uncertainties. The present calculation, however, represents the best information currently available. These results therefore provide a needed component for the present discussion on the potential impacts of these dams.

Comparison with Fossil Fuels

Comparisons without discounting

The annual emissions of greenhouse gases decline with time, but still stabilize at a level with significant impact. The timing of greenhouse-gas impacts, with emissions concentrated in the early years of a dam's life, is one of the principal differences between hydroelectric dams and fossil fuels in terms of global warming (Fearnside 1997b). Giving greater weight to short-term impacts increases the impact of hydroelectric dams relative to fossil fuels.

The fossil-fuel carbon displaced can be calculated based on the assumption that the alternative is generation from natural gas. This is a more reasonable assumption than oil as a baseline case, since current expansion of generating capacity for São Paulo and other parts of the electrical grid in Central-South Brazil is now coming from gas-fired plants that are fuelled from Brazil's new gas pipeline from Bolivia. The gas pipeline already exists and is not considered as part of the construction emissions of the gas-fired power plants.

Fossil-fuel displacement is shown in Figure 7 on an annual basis. The complex begins gaining ground on compensating for its emissions after year 15. The balance with greenhouse-gas emissions on a cumulative basis is shown in Figure 8. The complex only breaks even in terms of its global-warming impact 41 years after filling the first dam.

[Figures 7 and 8 here]

The longer the time horizon, the lower the average impact. For the first ten years the net impact is 4.0 times that of the fossil-fuel alternative; after twenty years the net impact is still 2.5 times that of fossil fuel, while for the full 50-year time horizon the project pays back its global-warming debt (assuming that it is interest free—that is, calculated with zero discount), with the long-term average gross impact at 70% that of the fossil-fuel alternative.

The effect of time

The role of time is an essential part of the debate on hydroelectric dams and on the question of global warming in general. Most decisions, such as a decision to build a

hydroelectric dam, are based on financial cost/benefit calculations that assign an explicit value to time by applying a discount rate to all future costs and benefits. The discount rate is essentially the opposite of an interest rate, such as one might earn on a savings account in a bank. With a savings account, the longer one waits the greater the monetary amount in the account, as the balance is multiplied by a given percentage at the end of each time period and the resulting interest is added to the balance for compounding in the next time period. With discounting the value attached to a future amount decreases, rather than increasing, by a fixed percentage in each time period. If a project such as a hydroelectric dam produces great impacts in the first years, such as the tremendous peak of greenhouse gas emissions shown here, while the benefits in the form of fossil-fuel substitution from power generation only accrue over the long term, then any positive discount rate will weigh against the hydropower option (Fearnside 1997b).

The timing of greenhouse-gas emissions further increases the dam's impact when the emissions of the cement, steel and fossil fuel used in the dam's construction are counted. The emissions from dam construction come years before any electricity is generated. A "Full Energy Chain," or FENCH, analysis would include all of these emissions. However, the construction emissions are a relatively small of the total impact. Discounted annual net emissions at rates up to 3% are shown in Figure 9. If only the instantaneous balance is considered, the complex substitutes for more carbon equivalent than it emits beginning in year 16, independent of discount rate. Thereafter the complex begins to pay off its environmental debt for the heavy net emissions in the first 15 years.

[Figure 9 here]

The discounted cumulative emissions peak in year 15, but do not reach the break-even point until at least 41 years after the first reservoir is filled (Figure 10). Discounting substantially lengthens the time needed to achieve this breakeven point.

[Figure 10 here]

The effect of different annual discount rates is shown in Figure 11. At zero discount the average net impact represents an annual gain of 1.4 million Mg C (the 50-year average in Table 9), but the relative impact attributed to hydropower increases greatly when the value of time is considered. In the case of the Belo Monte/Babaquara complex, any annual discount rate above 1.5% results in the project having a greater impact on global warming than the fossil-fuel alternative. Discount rates up to 12% are shown; although this author does not advocate such heavy discounting (Fearnside 2002b,c), one important contingent in the debates over carbon accounting (for example, the European Forestry Institute) advocates using the same discount rates for carbon as for money, and the financial analyses for Belo Monte use a 12% discount rate for money (Brazil, ELETRONORTE 2002, Tomo I, p. 6-84).

[Figure 11 here]

In terms of global warming, a series of arguments provide a rationale for assigning a value to time in greenhouse-gas emission calculations (Fearnside 1995b, 1997b, 2002b,c, Fearnside and others 2000). Global warming is not a one-time event

(such as a volcanic eruption or a tidal wave)—temperature increase creates an essentially permanent change in the probability of droughts and other environmental impacts. Any delay in greenhouse-gas emissions and consequent temperature increases therefore represents a savings of the human lives and other losses that would otherwise have occurred over the period of the delay. This gives time a value that is independent of any “selfish” perspective of the current generation. Despite the benefits of giving value to time in order to favor decisions that delay global warming, reaching political agreement on appropriate weights for time is extremely difficult. The decision in the first rounds of negotiations over the Kyoto Protocol has been to use a 100-year time horizon with no discounting over this period as the standard for comparing the different greenhouse gases (*i.e.*, the global-warming potential of 21 adopted for methane). If alternative formulations are used that give weight to time, the impact of the Belo Monte/Babaquara complex would increase, and, more importantly, it would increase the development’s impact relative to other possible options for energy supply.

Policy Implications

The finding of the present study that the proposed Belo Monte and Babaquara (Altamira) dams would produce substantial net emissions of greenhouse gases for many years is an important consideration for ongoing policy debates in Brazil and in other countries facing similar decisions. The additional greenhouse-gas emission of 11.2 million Mg of CO₂-equivalent carbon per year over the first ten years represents more than the current fossil-fuel emission from metropolitan São Paulo, which is home to 10% of Brazil’s population. Rational decision making on proposals for hydroelectric dams, as with any development project, requires a comprehensive assessment of both the impacts and the benefits of proposals so that the pros and cons can be compared and publically debated prior to making decisions on project construction. Greenhouse gases represent an impact that has so far received little consideration in these decisions.

In the case of Belo Monte and Babaquara (Altamira), it is important to recognize that the benefit side of the balance is considerably less attractive than is often portrayed by project proponents. The electricity produced is for a grid that supports a rapidly growing sector of subsidized electro-intensive industries, such as aluminum smelting for export. The aluminum sector in Brazil employs only 2.7 people per GWh of electricity consumed, second only to iron-alloy smelters (1.1 job/GWh), which also consume large amounts of energy for an export commodity (Bermann and Martins 2000, p. 90). A national discussion on the use that is made of the country’s electricity should be a prerequisite for a major decision to increase generating capacity, as by building the Xingu dams. The contrast between the social costs of dams and the meager benefits they provide through the electro-intensive industries they support is particularly relevant to the plans for the Xingu River (Bermann 2002, Fearnside 1999).

From the point-of-view of greenhouse gases, the fact that energy is used for a subsidized export industry means that the baseline against which hydroelectric emissions are compared should perhaps include simply not producing some of the power expected from the dams, rather than the baseline used here of generating the full equivalent of the dams’ power output from fossil fuels. Because Brazil could choose not to expand or maintain its electro-intensive export industries, such an alternative baseline would make the emissions results even less favorable for hydropower than those calculated in the present paper.

The Xingu River dams represent a challenge to Brazil's environmental licensing system because of the great difference in impact between the first dam (Belo Monte) and the subsequent dams, especially Babaquara (Altamira). Brazil's environmental licensing system currently only examines the impacts of one project at a time, not the combined impact of interdependent projects such as these. Because the greatest impacts (including greenhouse-gas emissions) of a decision to build Belo Monte would be caused by the dam or dams that would consequently be built upstream, the licensing system must be reformed to cope with this type of situation.

Conclusions

The Belo Monte and Babaquara (Altamira) dam complex would have a substantial impact on global warming, although the large amount of energy produced would eventually compensate for the high initial emissions. The assumptions used here indicate that 41 years would be necessary for the complex to break even in terms of global-warming impact if no discounting is applied. Despite high uncertainty in a number of key parameters, the general conclusion appears to be robust that the complex would have substantial impact, and that the long-term level of impact, although much lower than that in the first few years, would be maintained at appreciable amounts.

The present analysis includes number of conservative assumptions regarding percentages of methane emitted by different pathways. Higher values for these parameters would further extend the time needed for the complex to break even in terms of global-warming impact.

The impact attributed to dams is highly dependent on any value given to the timing of emissions: any form of discounting or other time-preference mechanism applied would further increase the dam's calculated impact relative to generation from fossil fuels. The value of 41 years for a greenhouse-gas emission of this magnitude is substantial even at zero discount. With annual discount rates above 1.5% the dams fail to break even by the end of the 50-year time horizon.

The case of Belo Monte and the other Xingu dams illustrates the absolute necessity of considering the interconnections among different infrastructure projects and including these considerations as a precondition for constructing or licensing any of the projects. Postponing analysis of the more controversial projects is not a solution.

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FIGURE LEGENDS

Figure 1. A.) Babaquara (Altamira) Reservoir; B.) Belo Monte Channel and Canals Reservoirs. Sources: Babaquara: Brazil, ELETRONORTE nd[C. 1988]; Belo Monte: Brazil, ELETRONORTE nd[C. 2002]a.

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Table 1: Technical characteristics of Belo Monte (Kararaô) and Babaquara (Altamira) Dams

Item	Units	Belo Monte (Kararaô)			Babaquara (Altamira)	Notes
		Channel reservoir	Canals reservoir	Belo Monte total		
Reservoir area at maximum normal operating level	km ²	333	107	440	6,140	
Reservoir area at minimum normal operating level	km ²	333	102	438	2,560	
Drawdown area	km ²	0	5	5	3,580	
Water level fluctuation	m	0	1		23	
Volume at maximum normal operating level	Billion m ³	2.07	1.89	3.96	143.5	
Volume at minimum normal operating level	Billion m ³	2.07	1.79	3.86	47.16	
Live storage volume	Billion m ³	0	0.11	0.11	96.34	
Average depth	m	6.2	17.7	9.0	23.4	
Average residence time	days	3.1	2.8	5.8	211.6	(a)
Reservoir length	km	60	87	147	564	
Shoreline length	km	361	268	629	2,413	(b)
Number of turbines	Number	7	20	27	18	
Maximum output per turbine	MW	25.9	550	--	348.6	
Installed capacity	MW	181.3	11,000	11,181.3	6,274	
Water consumption per turbine	m ³ /s	253	695	--	672	
Total water consumption	m ³ /s	1,771	13,900	15,671	12,096	
Average streamflow	m ³ /s			7,851	7,851	(c)
Elevations						
Normal maximum operating level	m above msl	97	97	--	165	
Normal minimum operating level	m above msl	97	96	--	142	(d)
Spillway level	m above msl	76	79.52	--	145	(e)
Adduction canal level	m above msl	--	84	--	--	
Turbine intake axis	m above msl	80	65	--	116.5	
Other parameters						
Drainage area	km ²			447,719		
Annual evaporation	mm			1,575		
Annual precipitation	mm			1,891		
Location						
Location	Latitude	03° 26' S	3° 7' 35" S		3° 18' 0" S	
	Longitude	51° 56' W	51° 46' 30" W		52° 12' 30" W	

Notes

(a) Assumes all water used by main powerhouse at Belo Monte.

(b) Babaquara assumed to have same shoreline to length ratio as Belo Monte.

(c) Streamflow is the 1931-2000 averaged used in the Belo Monte EIA. A higher "synthetic" streamflow of 8,041 m³/s was calculated by Maceira and Damázio (nd) for Babaquara.

(d) The Canals Reservoir will have an unusual water management, with the level at 96 m during the high-flow season and 97 m during the low-flow season (Brazil, ELETRONORTE 2002).

(e) Babaquara spillway assumed 20 m below normal operating level

Table 2: Ratio of CH₄ concentrations to concentration at 30 m depth

Depth range (m)	Age ≤ 12 months ^(a)	Age 12.1-36 months ^(b)	Age > 36 months ^(c)
0 – 0.9	0.33	0	0
1 – 1.9	0.50	0	0
2 – 4.9	0.75	0	0
5 – 9.9	0.83	0	0.34
10 – 14.9	0.67	0	0.63
15 – 19.9	0.75	0.33	0.71
20 – 24.9	0.83	0.50	0.79
25 – 29.9	0.92	0.83	0.89
30 – 30.0	1.00	1.00	1.00
≥ 31	(d)	(d)	(d)

(a) Data from the Samuel reservoir 5 months after filling (J. G. Tundisi, cited by Rosa and others 1997, p. 43).

(b) Data from Petit Saut (Galy-Lacaux and others 1999).

(c) Data from Tucuruí 44 months after filling (J. G. Tundisi, cited by Rosa and others 1997, p. 43).

(d) $Y = 1 + (0.0165 (X - 30))$

where: Y = Ratio of CH₄ concentration to the concentration at 30 m depth

X = Depth below the surface (m)

Table 3: Bubbling and diffusion emissions from CH₄ concentration at 30 m depth^(a)

Emission pathway	Depth range (m)	CH ₄ concentration range at 30 m depth (mg/liter)	Slope	Intercept
Bubbling	0-3 m	0-9.2	47.572	-54.214
		9.3-17.6	64.979	-216.344
		≥ 17.7	23.562	516.453
Bubbling	4-6 m	0-9.2	31.284	-77.499
		9.3-17.6	35.738	-118.989
		≥ 17.7	12.959	284.049
Bubbling	7-8 m	0-4.5	0	0
		4.6-9.2	2.468	43.680
		9.3-17.6	11.139	-37.087
		≥ 17.7	4.039	88.535
Diffusion	All depths	0-9.2	11.909	-35.860
		9.3-17.6	17.917	-91.822
		≥ 17.7	1.895	191.656

(a) $Y = m X + b$

where: $Y = \text{CH}_4 \text{ emission (mg/m}^2\text{/day)}$

$X = \text{CH}_4 \text{ concentration range at 30 m depth (mg/liter)}$

$m = \text{Slope}$

$b = \text{Intercept}$

Parameter	Value	Units	Source
Above-ground fraction	0.759		Fearnside 1997b, p. 3
Average depth of surface water zone	1	meter	Assumption, based on
Leaf decay rate in seasonally inundated zone	-0.5	Fraction/year	Assumption.
Above-water decay rate (0-4 yrs)	-0.1680	Fraction/year	Assumed same as fell
Above-water decay rate (5-7 yrs)	-0.1841	Fraction/year	Assumed same as fell
Above-water decay rate (8-10 yrs)	-0.0848	Fraction/year	Assumed same as fell
Above-water decay rate (>10 yrs)	-0.0987	Fraction/year	Assumed same as fell
Carbon content of wood	0.50		Fearnside and others Revilla Cardenas (198
Average total biomass of forest at Babaquara	244	Mg/ha	Above-ground fraction
Average water depth at minimum level	18.4	meters	At 142 m above msl
Average water depth at normal operating level	23.4	meters	At 165 m above msl
Initial biomass present: leaves	4.1	Mg/ha	Calculated from total b
Initial biomass present: wood above water	138.8	Mg/ha	Calculated from total b
Initial biomass present: below ground	58.8	Mg/ha	Calculated from total b
Methane release by termites in forest	0.687	kg CH ₄ /ha/year	Martius and others 199
Methane release by termites in above-water biomass per Mg C decayed through termites	0.0023	Mg CH ₄	Martius and others 199
Percent of decay mediated by termites above the normal operating level waterline	4.23	%	Martius and others 199
Percent of decay mediated by termites below the normal operating level waterline	0	%	Based on Walker and
Total area of reservoir at normal operating level	6,140	km ²	
Riverbed area	136	km ²	Revilla-Cardenas 1988
Area cleared prior to flooding (permanently flooded zone)	0	km ²	
Total area of forest flooded	6,004	km ²	Calculated by differenc
Original forest area of permanently flooded zone	2,424	km ²	Area of zone, minus ri
Original forest area of drawdown zone	3,580	km ²	Calculated by differenc
Rate of trunks breaking off at waterline for trees > 25 cm DBH	0.063	Fraction of original stock/year	Based on Walker and
Rate of trunks breaking off at waterline for trees < 25 cm DBH	0.113	Fraction of original stock/year	Based on Walker and othe
Rate of branch fall (and assumed fall of trunks above first branch)	0.094	Fraction of original stock/year	Based on Walker and othe
Percent of above-ground live wood biomass in branches and trunk above first branch	30.2	%	Fearnside 1995a, p. 12 ba

Bole as percent of total live above-ground biomass in live trees > 10 cm DBH	57.47	%	Based on biomass expansion; biomass > 190 Mg/ha in li
Above-ground live biomass < 10 cm DBH	22.2	Mg/ha	Calculated from above
Branches as percent of live bole biomass	51.4	%	Based on Brown and Lugo
Branch biomass	55.9	Mg/ha	Calculated from above
Above-ground biomass of forest at Babaquara	185.3	Mg/ha	Calculated from total and
Above-ground live wood biomass	155.5	Mg/ha	Total-leaves-dead
Above-ground dead wood biomass	25.6	Mg/ha	Klinge 1973, p. 179
Live bole biomass	108.6	Mg/ha	Partioning based on Brown
Live bole biomass 10-25 cm DBH	23.9	Mg/ha	Partioning based on Brown
Live bole biomass < 10 cm DBH	13.0	Mg/ha	Jordan and Uhl 1978
Live bole biomass 0-25 cm DBH	36.9	Mg/ha	Summed from above
Live bole biomass > 25 cm DBH	71.7	Mg/ha	Partioning based on Brown
Live bole biomass: above waterline	96.4	Mg/ha	Vertical distribution interpolation
Live bole biomass 0-25 cm DBH: above waterline	32.8	Mg/ha	Vertical distribution interpolation
Live bole biomass > 25 cm DBH: above waterline	63.6	Mg/ha	Vertical distribution interpolation
Fraction of trees that float	0.4	Fraction	Richard Bruce, personal communication
Fraction of original branches on remaining trees falling per year	0.094	Fraction	Calculated from Walker and
Average fraction of year drawdown area exposed	0.5	Fraction	Rough estimate based on

Vegetation type	BABAQUARA			BELO MONTE	
	Area (km ²)	Percent	Above- ground Biomass loading ^(b) (Mg/ha dry weight)	Area (km ²)	P
Unflooded vegetation					
Open terra firme forest [mixed open forest (FA) + open submontane forest (FS)]	3,565.3	58.0	175.2		
Open terra firme forest on steeply sloping terrain				205.7	
Open terra firme forest on rolling terrain				11.9	
Broadleafed secondary forest	10.9	0.2	20.0 ^(c)	11.0	
Flooded vegetation					
Dense seasonally flooded riparian forest or dense riparian forest (FC)	2,421.9	39.3	201.2	191.5	
Open seasonally flooded riparian forest (Alluvial pioneer grassy formations)	5.6	0.1	60.0		
No vegetation (river channel)					
Areas without plant cover	136.3	2.4	0.0	20.0	
Totals					
Total unflooded vegetation	3576.3	58.2		228.5	
Total flooded vegetation	2,427.5	39.4		191.5	
Total vegetation	6,003.7	97.6	185.3	420.0	
Total reservoir	6,140.0	100.0		440.0	

(a) Data from Revilla Cardenas (1987, p. 55; 1988, p. 87), with areas adjusted in proportion to the current reservoir are for Babaqura; 440 km² for Belo Monte).

(b) Values include dead biomass (litter and dead wood), vines, and root mat.

Item	Value	Units	Source
UPTAKE FROM UNFLOODED FOREST SOIL			
Average annual CH ₄ gas uptake from unflooded forest soil	-3.8	kg CH ₄ /ha/year	Potter and others (1996) from 22 studies
Total area of forest flooded by reservoir	6,004	km ²	Based on 6140 km ² reservoir area and streambed from Revilla Cardenas 1988, p. 87
Area of riparian forest flooded by reservoir	2,427	km ²	Revilla Cardenas 1988, p. 87
Area of <i>terra firme</i> forest flooded by reservoir	3,576	km ²	Calculated by difference
Fraction of year riparian forest flooded naturally	0.17	Fraction	Assumed 2 months on average
Uptake per ha per year in riparian forest	-3.17	kg CH ₄ /ha/year	Proportional to time unflooded
Uptake per year in riparian forest	-768.70	Mg CH ₄ /year	Uptake per ha × area of riparian forest
Uptake per year in <i>terra firme</i> forest	-1,358.98	Mg CH ₄ /year	Uptake per ha × area of <i>terra firme</i> forest
Total uptake per year	-2,127.68	Mg CH ₄ /year	Sum by forest type
Global warming potential (GWP) of CH ₄	21	Mg CO ₂ gas equivalent/ Mg CH ₄ gas	Schimel and others 1996
CO ₂ carbon equivalent/year	-0.012	Million Mg CO ₂ - equivalent C/year	Calculated from CH ₄ emission, GWP, atomic weight of C (12) and molecular weight of CO ₂ (44)

EMISSION BY FOREST TERMITES

Emission/ha/year	0.5	kg CH ₄ /ha/year	Fearnside 1996b
Ha-year equivalents of forest	0.6	Million ha-year equivalents	Calculated from above
Emission/year	317.0	Mg CH ₄ /year	Calculated from above
CO ₂ carbon equivalent/year	0.0018	Million Mg CO ₂ - equivalent C/year	Calculated as above

EMISSIONS FROM NATURAL FLOODING OF PRE-DAM FLOODED FOREST

Methane emission of flooded forest during natural flooding	103.8	mg CH ₄ /m ² /day	Mean of five studies in white-water <i>várzea</i> fore (Wassmann and Martius 1997, p. 140)
Days flooded per year	59.4	Days	Assumption, as above
Annual emission per km ²	6.2	Mg CH ₄ /year/km ²	Calculated from above
Annual natural emission by flooded forest	14,961	Mg CH ₄ /year	Calculated from above
CO ₂ carbon equivalent/year	0.086	Million Mg CO ₂ - equivalent C/year	Calculated from above
Annual annual emission adjusted for cycle length	44,883	Mg CH ₄ /year	Considering emission per cycle (2 months vs 6 months)
CO ₂ carbon equivalent/year	0.257	Million Mg CO ₂ -equivalent C/year	Calculated from above

EMISSIONS FROM PERIODIC PONDING EVENTS IN *TERRA FIRME* FOREST

Emission when flooded or ponded	103.8 mg CH ₄ /m ² /day	Assumed same as <i>várzea</i> forest flooding (as above).
Annual natural emission by ponding	187.0 Mg CH ₄ /year	Calculated from above
CO ₂ carbon equivalent/year	0.001 Million Mg CO ₂ -equivalent C/year	Calculated from above
TOTALS		
Total emission of methane	43,259 Mg CH ₄ /year	Calculated from above, including cycle-length adjustment
CO ₂ carbon equivalent/year	0.248 Million Mg CO ₂ -equivalent C/year	Calculated from above

Average annual N ₂ O gas emission from unflooded forest soil	8.7	kg N ₂ O/ha/year	Verchot and others 1999, p. 37
Total area of forest flooded by reservoir	6,004	km ²	Based on 6140 km ² reservoir area and streambed from Revilla Cardenas, 1988, p.87
Area of riparian forest flooded by reservoir	2,427	km ²	Revilla Cardenas 1988, p. 87
Area of <i>terra firme</i> forest flooded by reservoir	3,576	km ²	Calculated by difference
Fraction of year riparian forest flooded naturally	0.17	Fraction	Assumed 2 months on average
Emission per ha per year in riparian forest	7.23	kg N ₂ O/ha/year	Proportional to time unflooded
Emission per year in riparian forest	1,755.6	Mg N ₂ O/year	Emission per ha × area of riparian forest
Emission per year in terra firme forest	3,103.7	Mg N ₂ O/year	Emission per ha × area of <i>terra firme</i> forest
Total emission per year	4,859.3	Mg N ₂ O/year	Sum by forest type Schimel and others 1996
Global warming potential (GWP) of N ₂ O	310	Mg CO ₂ gas equivalent/ Mg N ₂ O gas	
CO ₂ carbon equivalent/year	0.411	Million Mg CO ₂ - equivalent C/year	Calculated from area and per-hectare emission, atomic weight of C (12) and molecular weight of CO ₂ (44)

EMISSIONS FROM FLOODING

Riparian forest flooding	404.6	km ² -days/year	Calculated from area and assumed 6-month average flood
<i>Terra firme</i> forest ponding	1,801.1	km ² -days/year	Calculated from area, 5% flooding per event (based on Mori and Becker 1991) and assumed 5-year frequency and 30 day duration
Emission when flooded	7.6	kg N ₂ O/km ² -day	7.6 mg N ₂ O/m ² /day (Tucuruí and Samuel reservoir means from de Lima and others 2002)
Emission from riparian forest flooding	3.1	Mg N ₂ O/year	Flooding (km ² /days) × emission/km ² /day
Emission from <i>terra firme</i> forest ponding	13.7	Mg N ₂ O/year	Ponding (km ² -days) × emission/km ² /day
Total emission from flooding	16.8	Mg N ₂ O/year	sum of flooding emission
CO ₂ carbon equivalent/year	0.001	Million Mg CO ₂ - equivalent C/year	Calculated from GWP as above.

TOTAL EMISSION

Total emission from forest loss	4,876.0	Mg N ₂ O/year	Sum of unflooded forest, ponding and flooding emissions
CO ₂ carbon equivalent/year	0.412	Million Mg CO ₂ - equivalent C/year	Calculated from GWP as above.

Table 8: Greenhouse-gas emissions from dam and transmission-line construction

Item	Units	Emission per unit CO ₂ -equiv C (kg)	Ref- erence	Note	Belo Monte		Emission (million Mg CO ₂ - equivalent C)	Note	Babaquara	
					No. of units	Ref- erence			No. of units	
Steel	Mg	600.0	(a)	(b)	323,333	(c)	0.194	(a)	30	
Cement	kg	0.207	(a)	(c)	848,666,000	(e)	0.176	(a)	1,217,25	
Diesel	million kg	863,280	(d)		135.1	(f)	0.117			
Electricity	TWh	139,903,200	(g)		3.15	(h)	0.441			
Dam sub-total							0.928			
Transmission line construction	km-MW installed	1.9	(i)		29,596,901	(j)	0.055	(a)	17,04	
Project total							0.983			

(a) Van Vate 1995.

(b) Uses IPCC 1994 100-year GWPs: CO₂=1, CH₄=24.5, N₂O=320 (Albritton and others 1995).

(c) Based on proportions relative to Belo Monte (see text).

(d) Uses IPCC 1995 100-year GWP values [Kyoto Protocol values]: CO₂=1, CH₄=21, N₂O=310 (Schimel and others 1996).

(e) Based on concrete volumes (da Cruz 1996, p. 18)(see text).

(f) Babaquara diesel use based proportional to planned excavation.

(g) Based on displacement of combined-cycle gas in São Paulo (see text).

(h) Construction electricity use based on 280 kWh electricity/TJ (Dones and Gantner 1996). Emissions from electricity generation baseline in São Paulo (see text).

(i) Mean in Québec, Canada (Peisajovich and others 1996).

(j) Transmission lines from Belo Monte to South-Central Brazil branch and go to three destinations with a mean distance: Cachoeira Paulista-SP (2662 km), Campinas-SP (2599 km) and Ouro Preto-MG (2680 km) (Brazil, MME-CCPESE 2006). Babaquara has an additional 70 km of line.

Table 9: Long-term averages of net greenhouse-gas emission for the Belo Monte/Babaquara complex

Emissions from all sources (million Mg CO ₂ -equivalent C/year)			
	Years 1-10 10-year average	Years 1-20 20-year average	Years 1-50 50-year average
Inventory emissions			
Surface emissions	1.0	0.8	0.4
Turbines	2.6	3.8	2.8
Spillway	1.6	1.0	0.6
Transposition canal	0.2	0.4	0.3
Bottlenecks	0.01	0.01	0.01
Total inventory emissions	5.3	6.0	4.1
Other fluxes			
Fossil-fuel substitution	-3.7	-3.9	-4.1
Pre-dam ecosystem fluxes	-0.3	-0.5	-0.6
Above-water biomass	9.6	7.2	3.8
Shoreline decay	0.07	0.04	0.01
Total other fluxes	5.9	0.1	-5.5
Total impact	11.2	6.1	-1.4
Gross impact as multiple of fossil-fuel baseline	4.0	2.5	0.7

Table 10: Sensitivity tests for selected input parameters

Input parameter	Units	Reference scenario parameter value	Percent change in total impact in response to a 10% increase in input parameter		
			10-year average	20-year average	50-year average
Proportion of CH ₄ released at turbines (midpoint of high and low scenarios)	Proportion	0.55			
Percent change	%		2.03	5.18	-16.17
Proportion of CH ₄ released at spillway	Proportion	1			
Percent change			1.22	1.37	-3.67
Growth rate of drawdown zone vegetation	Mg/ha/year dry weight	10			
Percent change	%		0.02	0.46	-3.33
Biomass of macrophytes, years 1-6	Mg/ha of macrophytes dry weight	11.1			
Percent change	%		0.23	0.26	-0.54
Biomass of macrophytes, years 7-50	Mg/ha of macrophytes	1.5			
Percent change	%		0.003	0.10	-0.48
Unbeached macrophytes (death+decay)	fraction per month	0.144			
Percent change	%		0.23	0.34	-0.92
Flooded drawdown (underwater decay rate)	fraction per month	0.656			
Percent change	%		0.02	0.48	-0.04
Beached macrophyte decay rate	fraction per month	0.314			
Percent change	%		-0.001	-0.01	0.04
Above-ground biomass	Mg/ha dry weight	176.1			
Percent change	%		2.95	5.02	-11.91
Adduction canal emission percent	%	15.3			
Percent change	%		0.02	0.08	-0.26
Adduction canal oxidation percent	%	66.5			
Percent change	%		-0.50	-2.49	7.90
Bottlenecks emission percent	%	2.1			
Percent change	%		0.01	0.004	0.002
Bottlenecks oxidation percent	%	9.2			

Percent change	%		-0.01	-0.07	-0.04
Shoreline mortality percent: 0-100 m (weighted average)	%	81.7			
Percent change	%		0.06	0.07	-0.04

Fig 1-a

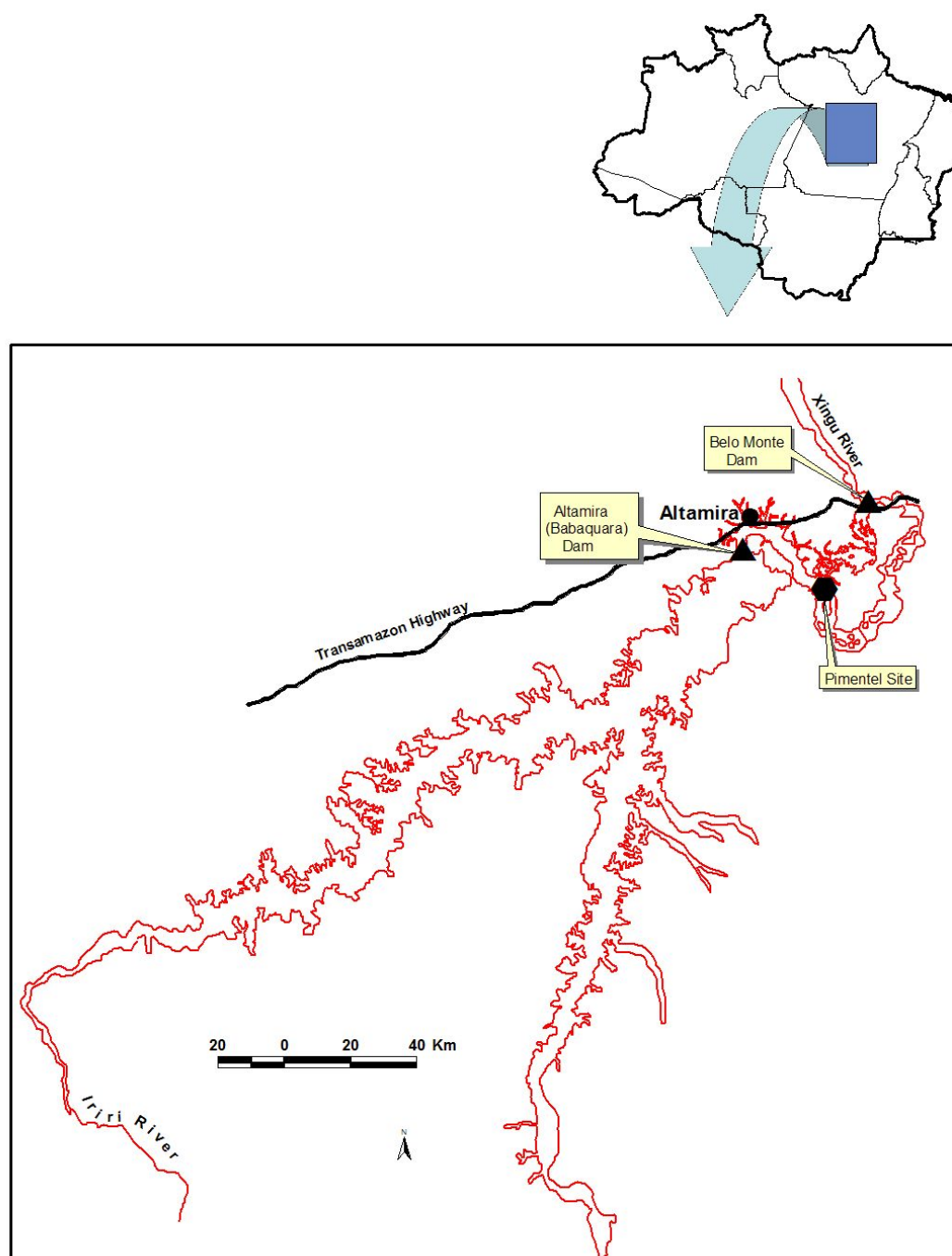


Fig 1-b

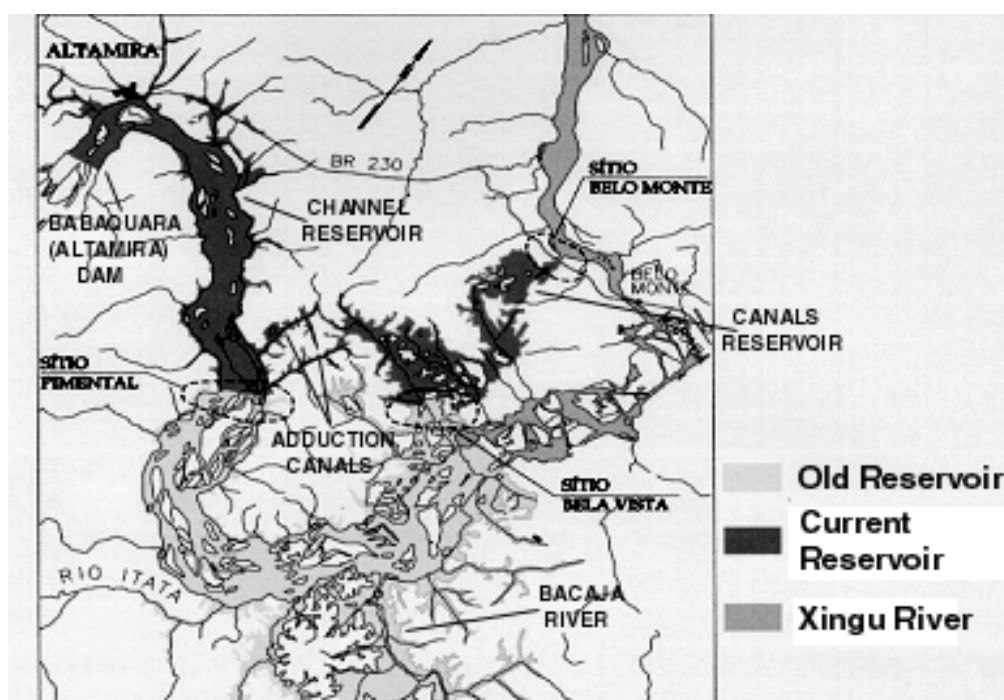


Fig. 2

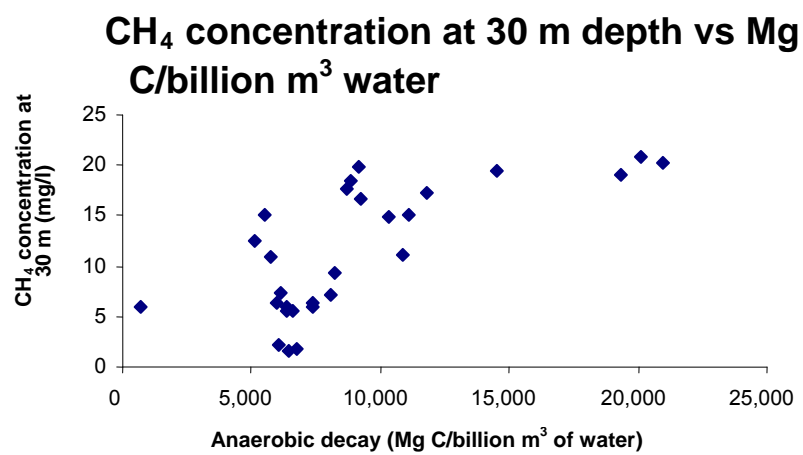


Fig. 3a.

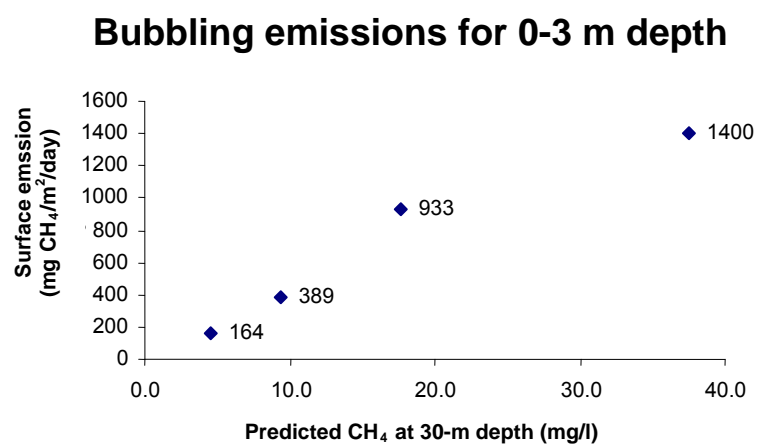


Fig 3b.

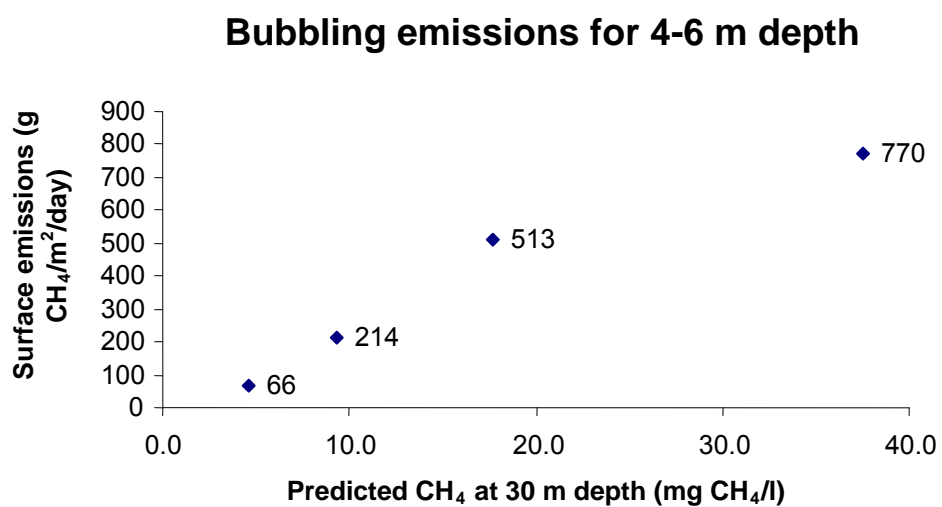


Fig. 3c.

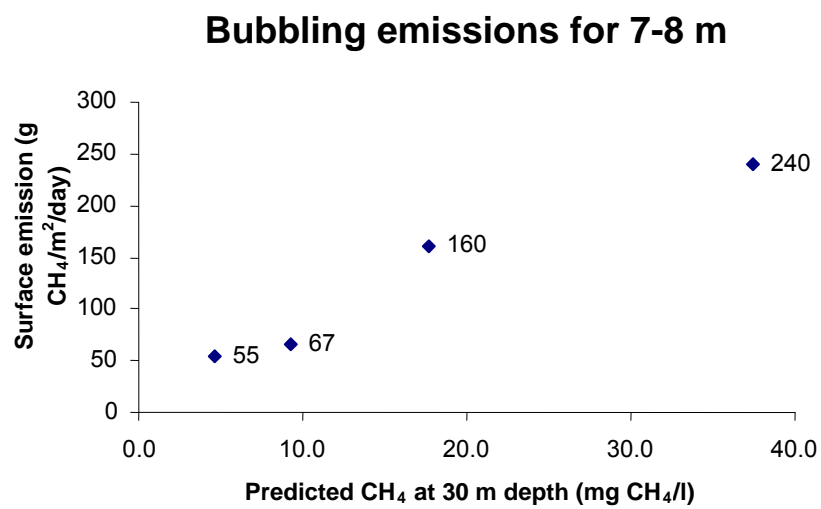


Fig. 3 d.

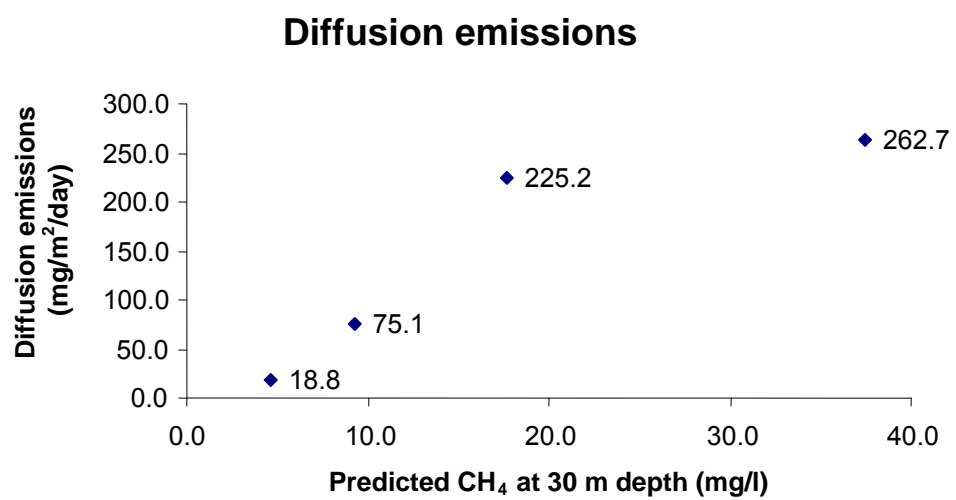


Fig. 4a

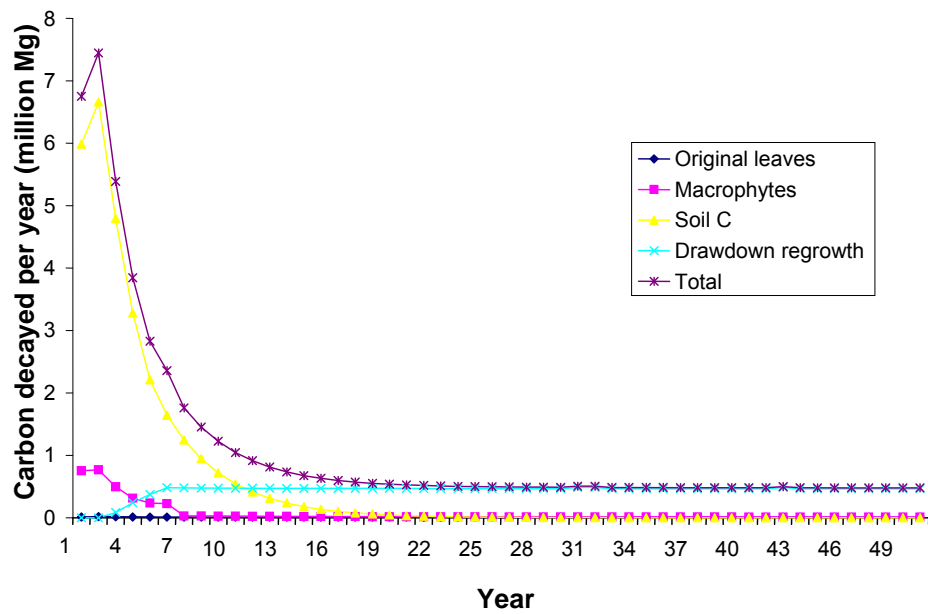
Babaquara: Carbon decayed anaerobically

Fig. 4b

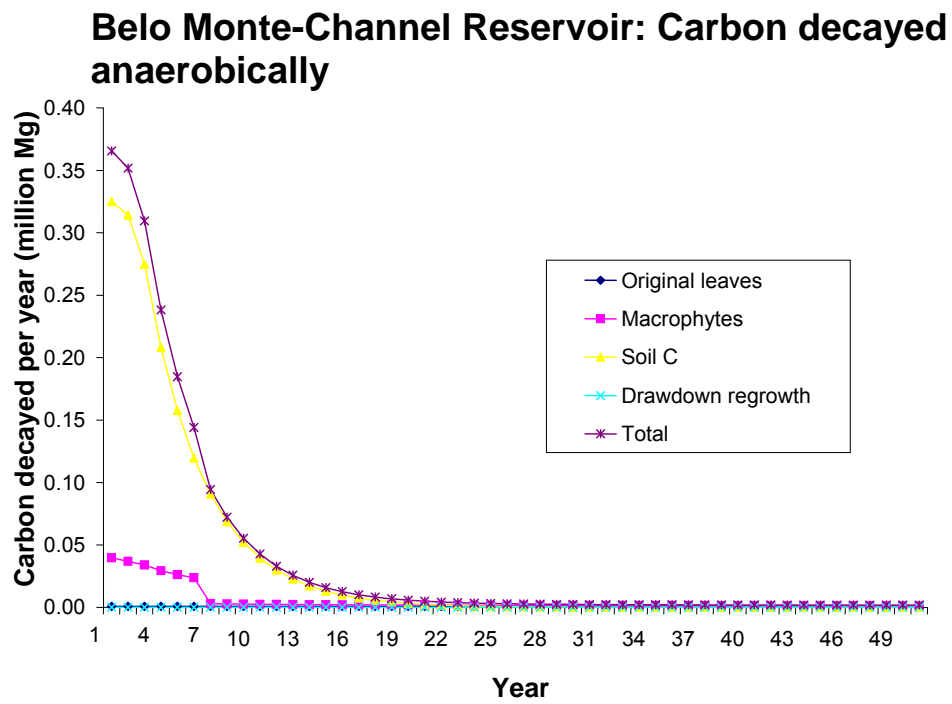


Fig. 4c

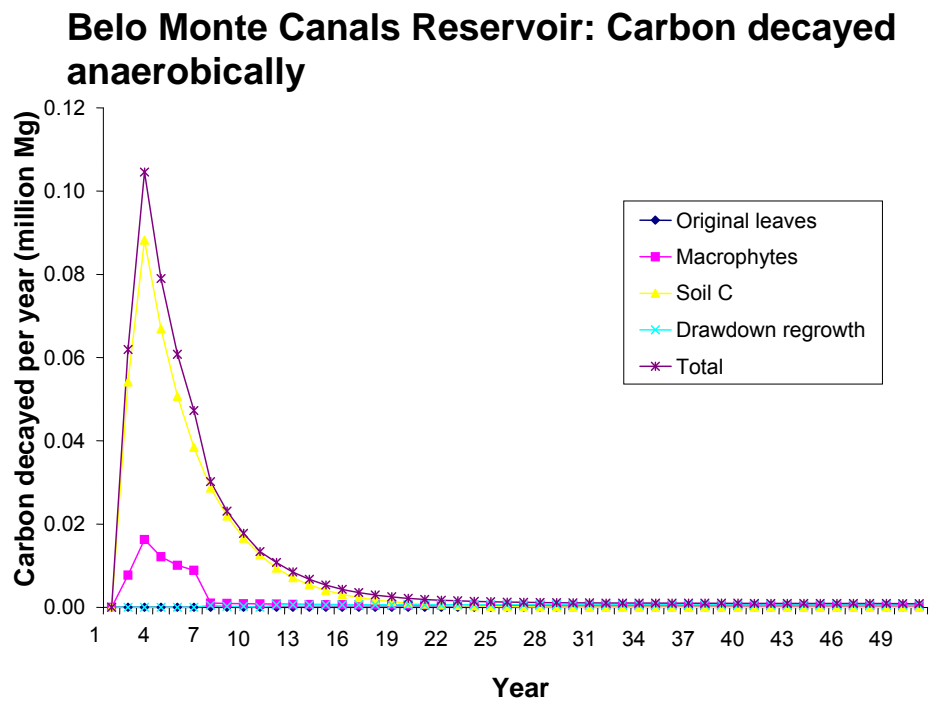


Fig. 5a

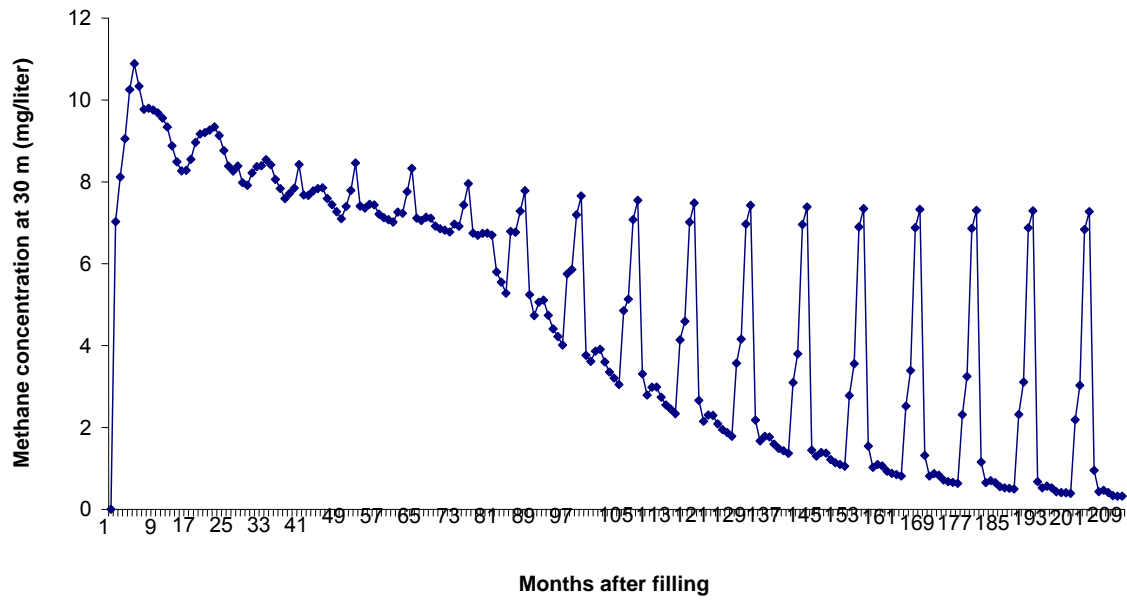
Babaquara: Calculated CH_4 concentration at 30 m depth

Fig. 5b

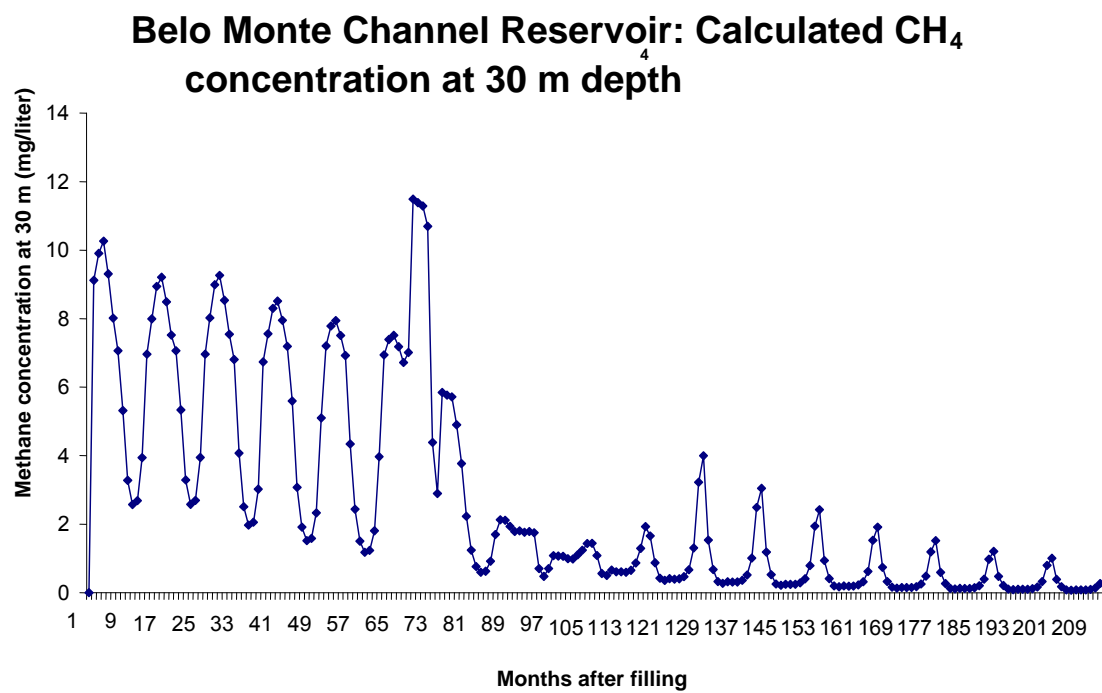


Fig. 5c

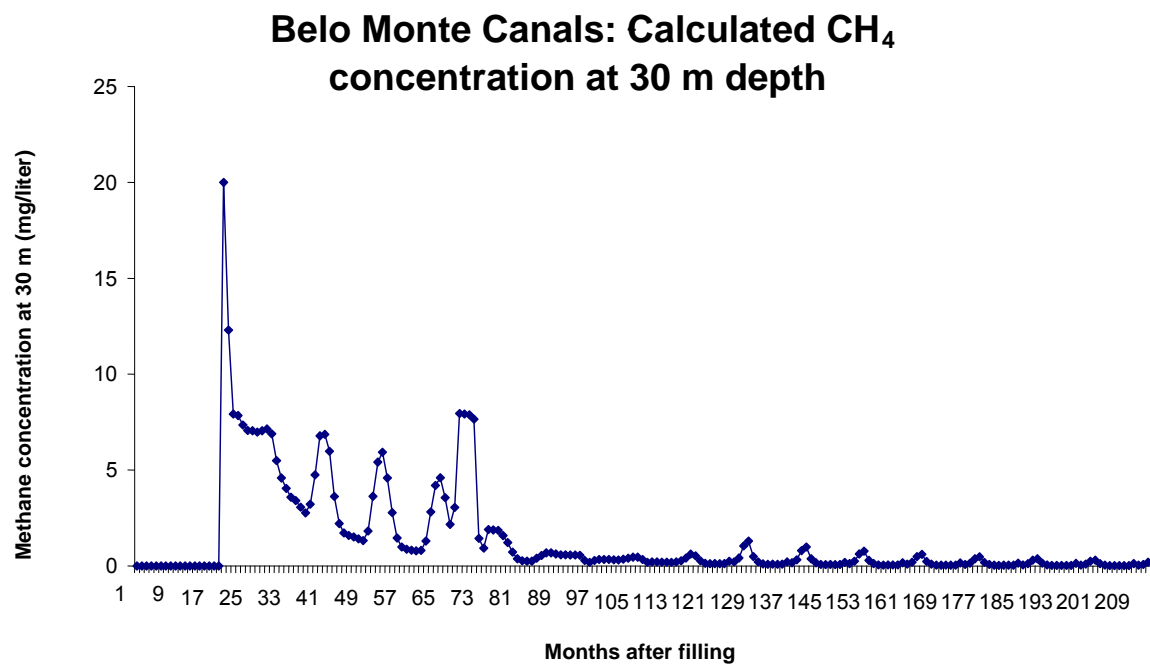


Fig. 6

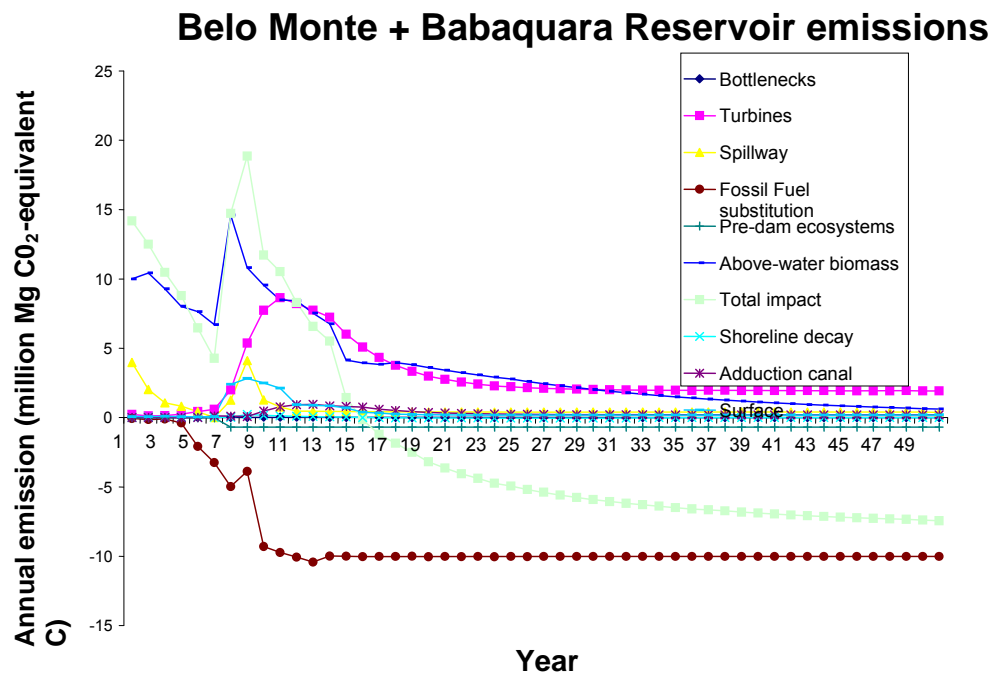


Fig. 7

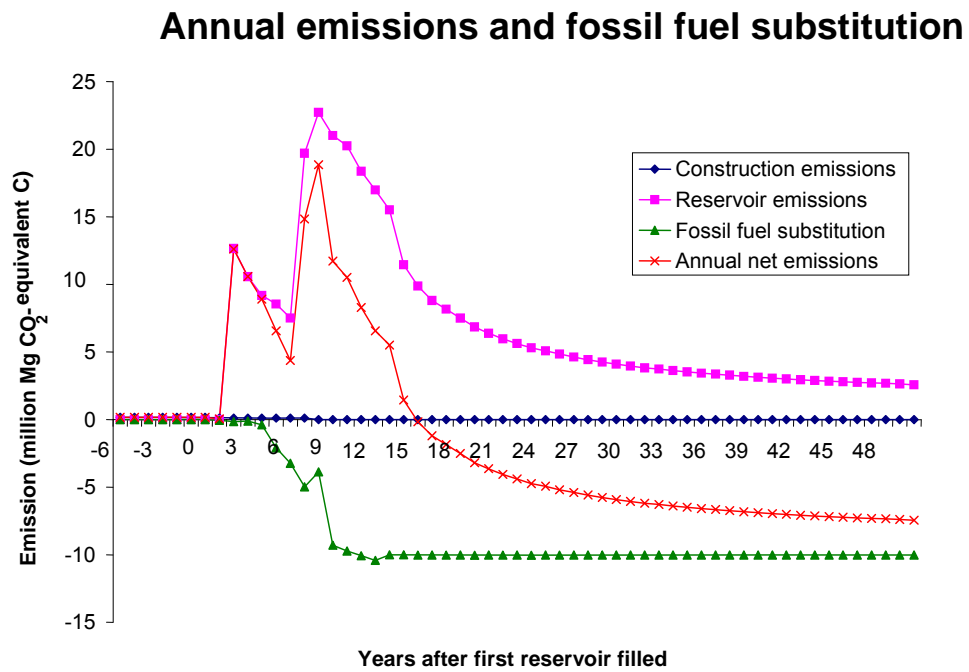


Fig. 8

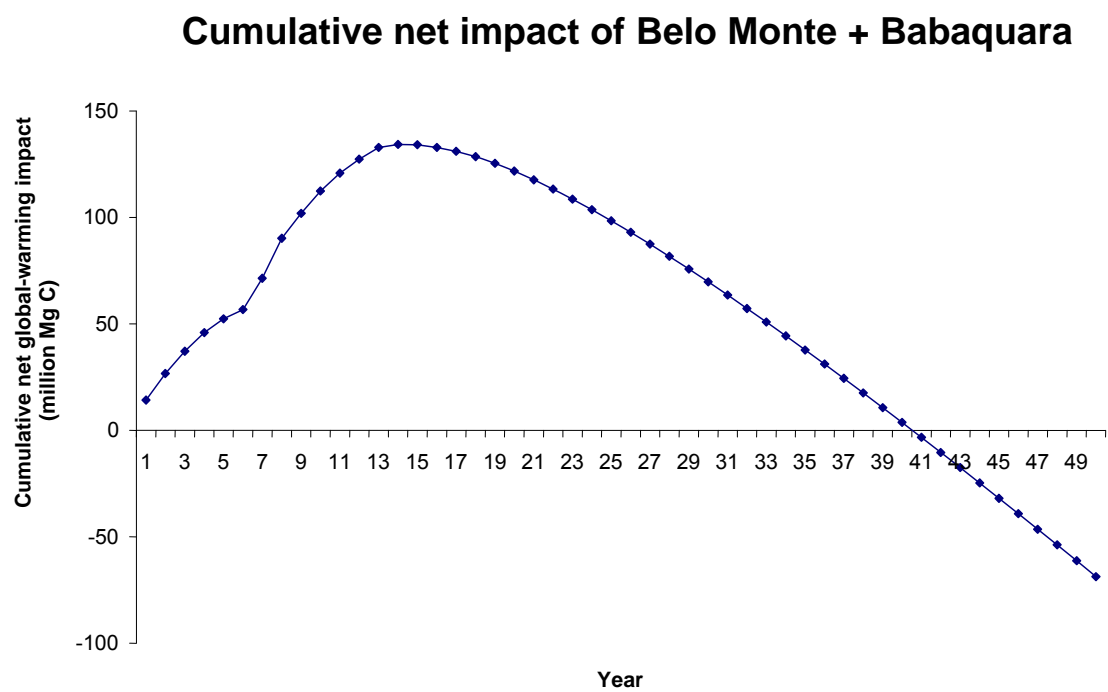


Fig. 9

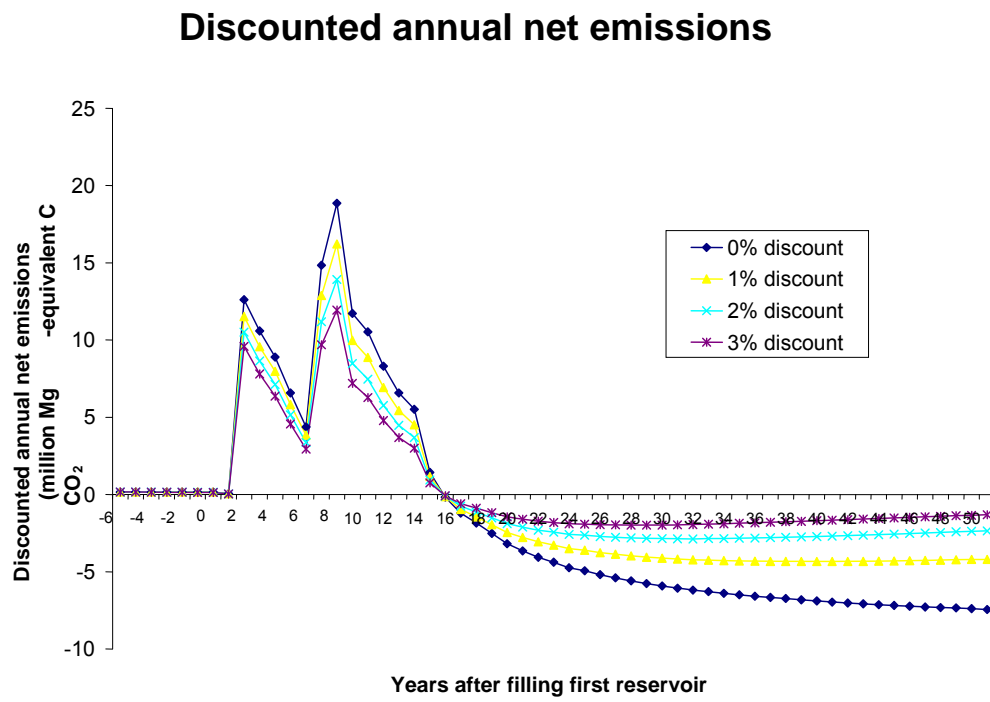


Fig. 10

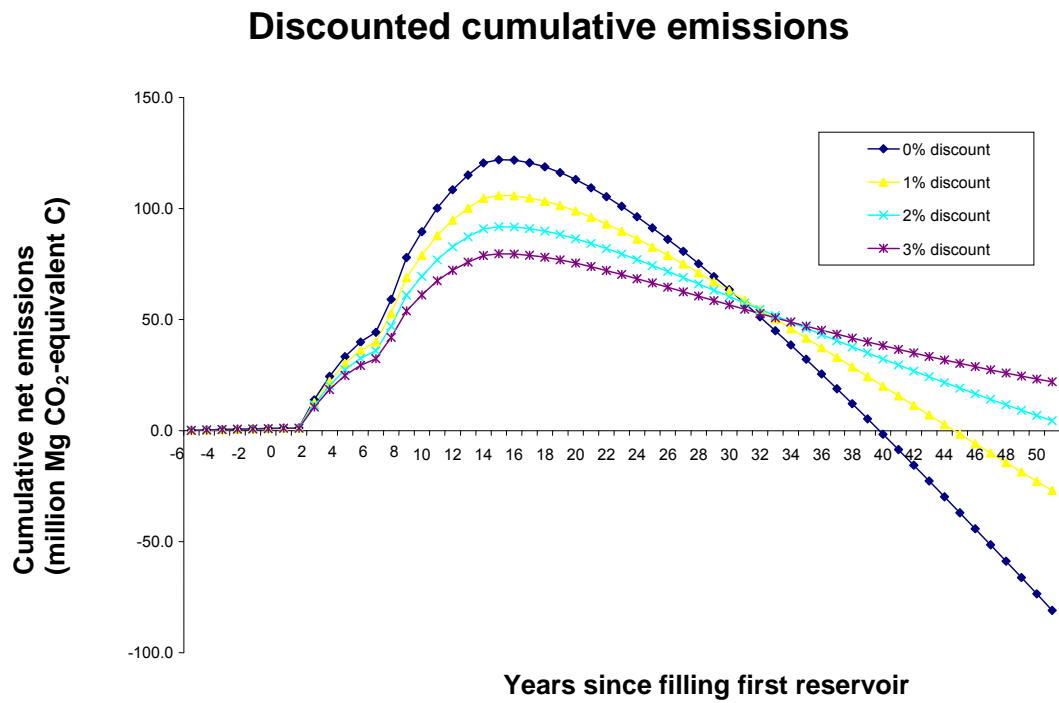


Fig. 11

